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**SHOCK TEMPERATURES IN SILICA GLASS:
IMPLICATIONS FOR MODES OF SHOCK INDUCED
DEFORMATION, PHASE TRANSFORMATION, AND
MELTING WITH PRESSURE**

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ABSTRACT

Greybody temperatures and emittances of silica glass under shock compression between 10 and 30 GPa are determined. Observed radiative temperatures are higher than computed continuum temperatures for shock compressed silica glass; however, below ~ 26 GPa observed emittances are < 0.02 . This suggests that fused quartz deforms heterogeneously in this shock pressure range as has been observed in other minerals. Between 10 and 16 GPa, radiative temperatures decrease from 4400 K to 3200 K, whereas above 16 GPa to 30 GPa greybody temperatures of ~ 3000 K with low emittances are observed. The emittances increase with pressure from 0.02 to 0.9. The pressure range from 10 to 16 GPa coincides with the permanent densification region while the 16 to 30 GPa range coincides with the inferred mixed phase region along the silica glass Hugoniot. The differing radiative behaviors may relate to these modes of deformation. Based upon earlier shock recovery experiments and a proposed model of heterogeneous deformation under shock compression, the temperatures associated with low emittances in the mixed phase region probably represent the melting temperature of the high pressure phase, stishovite, which can be expected to crystallize from a melt in hot zones. Above 20 GPa to 30 GPa the melting temperature of stishovite would therefore be $3000 \text{ K} \pm 200 \text{ K}$ and almost independent of pressure. The effect of pressure on melting relations for the system $\text{SiO}_2\text{-Mg}_2\text{SiO}_4$ are considered together with the proposed stishovite melting curve and suggest maximum solidus temperatures within the mantle of $\sim 2370 \text{ K}$ at 12.5 GPa and $\sim 2530 \text{ K}$ at 20.0 GPa. Using the proposed stishovite melting

temperatures T_m and estimates of upper mantle temperatures T , the effective viscosity, which can be considered a function of the homologous temperature T/T_m , appears to remain nearly constant from 600 to 200 km depth in the Earth.

INTRODUCTION

The temperatures of silica glass subject to shock pressures between 10 and 30 GPa were investigated by observation of shock induced radiative thermal emissions. In this pressure range many brittle materials deform heterogeneously under shock compression with the formation of localized zones of intense deformation and high temperatures surrounded by relatively undamaged cooler material. The study of these zones which are referred to as hot spots, shear bands, or microfaults is important because they play a role in the shock deformation process in minerals and particularly in shock production of glassy and high pressure polymorphs. Spectra of light emitted from silica glass in this pressure range reveal the temperature, and to some extent, the density of these microfaults. Observed temperatures appear to be related to the deformation occurring and, based on models of heterogeneous deformation, constrain the melting curve of minerals with pressure.

Quartz was chosen as a candidate for study for two major reasons. The first is the importance of quartz as both a crustal and a mantle mineral. High pressure phases of quartz have been discovered at sites of large meteorite impacts and are presumably produced by the action of shock waves. Knowledge of the shock effects in quartz provide information for deducing

the effects of large impacts on the Earth. Free SiO_2 in the form of coesite or stishovite most likely does not exist in substantial quantities in the mantle despite cosmochemical estimates that it composes $\sim 35\%$ by weight of the entire Earth, but is a constituent of the high pressure phases of pyroxene, olivine, and garnet. However, since the lower mantle is often described in terms of high pressure phase oxide mixtures [Al'tshuler *et al.*, 1965], a better understanding of the high pressure properties of SiO_2 yields information pertinent to the study of mantle constitution and dynamics.

Secondly, SiO_2 has been extensively studied in both high pressure static and dynamic experiments. The high pressure polymorphs coesite [Coes, 1953] and stishovite [Stishov and Popova, 1961] have room pressure densities of 2.92 Mg/m^3 and 4.29 Mg/m^3 [Kuskov and Fabrichnaya, 1987], respectively. Except for the discovery of coesite in kimberlites and deep-seated metamorphic rocks, the only natural occurrences of these minerals are at the sites of major impacts on the Earth's surface, suggesting formation from quartz by shock wave. Evidence [reviewed by Stöffler, 1972, 1974] demonstrate that shock damage in quartz is nonuniformly distributed; shock compression in quartz below 35 GPa is a heterogeneous process. Observation of the radiative temperatures of fused SiO_2 under shock compression will aid in further understanding of shock processes in brittle minerals.

Previous high pressure work on SiO_2 is reviewed and a model of the shock deformation process in silica glass is presented. Results of shock temperature measurements from fused quartz in the pressure range from 10 to 35

GPa are interpreted in light of this model and implications for the mantle are suggested.

BACKGROUND

We consider the microscopic effects of the passage of a shock wave through a solid brittle medium with emphasis on silica glass. Covalently bonded minerals of geophysical interest undergo brittle deformation according to *Grady* [1980]. Shock deformation of partially covalently minerals, such as SiO_2 , MgO , Al_2O_3 , and CaCO_3 , is characterized by all or some of the following features: 1) the existence of large Hugoniot elastic limits, 2) fluid-like release from high pressure, 3) nonequilibrium Hugoniot in the mixed phase region for minerals with a primary co-ordination phase change, and 4) the observation of heterogeneities in minerals recovered from a shock wave experiment.

The Hugoniot elastic limit (HEL) is the shock pressure at which dynamic yielding occurs; shock compression below the HEL is an elastic process, plastic deformation occurs above the HEL [*Davison and Graham*, 1979]. This results in a first shock propagating slightly faster than the longitudinal sound speed (elastic precursor) with magnitude equal to the HEL followed by a slower second shock described as a plastic wave with amplitude of the final shock pressure. Two wave shock structures have been reported in shock compression experiments on both alpha and fused quartz to 75 GPa [*Wackerle*, 1962; *Fowles*, 1967; *Sugiura*

et al., 1981]. They attribute the two-wave structure, observed over a wide pressure range, to anomalously high HEL in quartz which has been measured by a variety of methods to range from 5.5 to 15 GPa for differing crystal orientations [Fowles, 1967; Graham, 1974]. Similarly, estimates of the fused quartz HEL range from 8.8 to 9.8 GPa [Wackerle, 1962; Barker and Hollenbach, 1970; Sugiura *et al.*, 1981].

In contrast to metals which retain a finite shear strength above the HEL, evidence was presented in the above studies indicating that silica loses its rigidity above the elastic limit, i.e., the Hugoniot of quartz collapses to predicted static compression curves. Silica compressed above the elastic limit may be characterized as fluid-like [Graham, 1974]. This fluid-like behavior is best observed during the release of a shock compressed quartz crystal from high pressure. Grady *et al.*, [1975] measured sound velocities from rarefaction wave profiles on the Hugoniot for a quartz rock (novaculite).

The Hugoniot of α -quartz (Figure 1, filled symbols) supplies evidence for a phase transition, believed to be that of α -quartz to stishovite. In the shock wave velocity-particle wave velocity plot of Figure 1b, shock velocities remain constant over a range in U_p of 1.4 to 2.25 km/s. Although Wackerle [1962] and Fowles [1967] reported data for both the elastic precursor and the plastic shock, Figure 1 shows results for only the second or plastic wave. Trunin *et al.*, [1971] supplies information on one arrival only which may lead to the discrepancy between the data sets at $U_p < 1.4$ km/s. All data sets display a region of constant shock velocity characteristic of a phase change in alpha quartz as labelled

in Figure 1. This range of particle velocity is equivalent to the pressure range of 14.4 GPa to approximately 38 GPa and is hereafter referred to as the mixed phase region of the Hugoniot of alpha quartz. Within this range the incompressibility of quartz increases dramatically as is seen in Figure 1a of shock pressure vs. specific volume.

Wackerle [1962] first suggested that the produced high pressure phase was stishovite, or a like-co-ordinated liquid. Based on the comparable large density increase found in both the static and shock wave experiments, and on the fact that the shock wave and static pressures of phase initiation were similar (approximately 14 GPa) *McQueen et al.*, [1963] suggested shock production of stishovite occurred. They further suggested that the short time scale of the shock experiments prohibited large stishovite crystal growth and that instead a fluid with a short order stishovite-like structure may be created.

Between the HEL (~ 8.8 to 9.8 GPa) and approximately 16.0 GPa, fused silica undergoes a permanent densification as evidenced by shock recovery [Wackerle, 1962; Arndt et al., 1971] and particle velocity time profiles [Sugiura et al., 1981]. Gibbons and Ahrens [1971], however, find increased refractive index, indicative of permanent densification, in recovered fused quartz shocked to pressures as high as 30 GPa. This permanent densification has been more fully studied in static compression experiments on fused quartz [Bridgman and Simon, 1953] and is attributed to a rearrangement and tilting of the SiO_2 tetrahedra [Cohen and Roy, 1961; MacKenzie, 1963; Arndt and Stöfler, 1969; Grimsditch, 1984; and Hemley et al., 1986]. This densification is

observed to reverse upon annealing of both shock and statically compressed fused quartz to pressures in excess of 10 GPa.

Above 16 GPa the Hugoniot of fused quartz passes through a region in which the material may be in a mixed phase with stishovite. The inherent properties of shocked fused quartz have not yet been determined. Note that at approximately 30 GPa and above the Hugoniots of both crystalline and fused quartz converge suggesting that both are at the same phase state at higher shock pressures. This suggests the fused quartz mixed phase region may contain some stishovite. *Sugiura et al.*, [1981] suggest compression above 16 GPa is that of a high-density silica glass, they find little evidence for production of stishovite below 22 GPa. However, in recent reshock and release experiments on fused quartz [*Chhabildas and Grady*, 1984a] at 13 GPa and 21 GPa, the stishovite transformation was found to initiate at 13 GPa and to progress to a more fully mixed phase state at 20 GPa.

After shock compression a rarefaction or release wave provides a thermodynamic path upon isentropic decompression. The release wave path in the pressure-volume plane yields estimates of the zero pressure density of the high pressure phase. Experiments on crystal quartz [*Ahrens and Rosenberg*, 1968; *Grady et al.*, 1974; *Chhabildas and Grady*, 1984b; *Chhabildas and Miller*, 1985; and *Podurets et al.*, 1976] indicate: 1) The transformation to the high pressure phase occurs within the shock front, little further transformation occurs during the remainder of the compression, 2) the release path for quartz shocked to less than 10 to 15 GPa returns to the initial density, 3) the

release path of quartz shocked into the mixed-phase region (15 to 34 GPa) follows a path indicative of mixed quartz-stishovite, the ratio of stishovite increasing proportionally with pressure, and upon release to less than 8 GPa, reversion to a lower density approximately that of crystal quartz, and 4) above 50 GPa the release path is essentially that of the high pressure phase until a released pressure of 5 to 8 GPa is reached whereupon a reversion to a lower pressure phase with a density of $\sim 2.2 \text{ Mg/m}^3$ occurs.

The most direct evidence for the transformation to the high pressure polymorphs are the recovered phases from both natural impact craters and laboratory shock experiments. Abundant coesite and smaller amounts of stishovite have been detected in rocks subjected to high shock pressures during meteorite impacts [*Chao*, 1960,1962; *Kieffer*, 1971]. Shock laboratory production of stishovite from quartz in single crystals, novaculite, and sandstone was reported by *DeCarli and Milton* [1965]. The lack of coesite in their recovered samples is not understood. *Kleeman and Ahrens* [1973] pointed out that the direct transformation from quartz to stishovite should occur at a pressure of ~ 7.5 GPa and not 14 GPa and identified stishovite from quartz shocked above 9.0 GPa; increasing quantity was found with increasing pressure until at pressures above 23.4 GPa the production of stishovite decreased with more glass produced. A single shot with fused quartz at 16.6 GPa yielded the largest percentage of stishovite. Most recently, *Ashworth and Schneider* [1985] report bands of polycrystalline stishovite, producing rings in electron diffraction, associated with densified glass in shock

recovered quartz.

Mashimo et al., [1980] suggest that stishovite is not readily observed in shock recovered material due to its rapid annealing to fused quartz at the elevated residual temperatures upon shock release. Above 600 K stishovite readily transforms to fused quartz at standard pressure [*Dachille et al.*, 1963].

MODEL OF HETEROGENEOUS DEFORMATION IN QUARTZ

Anan'in et al., [1974a] examined recovered crystal quartz and found the material in a plane perpendicular to the direction of shock propagation was broken up into 5-20 micron blocks separated by layers of quartz glass. The plane parallel to the shock front revealed larger (100-200 micron) blocks. *Anan'in et al.*, [1974a] suggested that compression above the HEL occurs by fragmentation into blocks which then move relative to one another. A consequence of this movement is frictional heating and melting at the surfaces of the blocks. They further speculated that increases in atomic co-ordination can occur rapidly within the melt and therefore the melted component, which increases with increasing pressure, accounts for the mixed phase region of the quartz Hugoniot.

Anan'in et al., [1974b] in shock recovery experiments on fused quartz found that it undergoes deformation similar to that of crystal-line quartz. As in crystal quartz, they note that at shock pressures above the HEL, the fused quartz splits into blocks separated by layers of melt. The solid blocks are thus subjected to hydrostatic compression and are permanently densified whereas the melt takes on the characteristics of a

high pressure phase. Moreover, *Wackerle* [1962] found that fused quartz shocked to 25 GPa transformed to fine particles with densities in the 2.4-2.5 Mg/m^3 range whereas material shocked to higher pressures had the normal 2.2 Mg/m^3 density in agreement with index of refraction data [*Gibbons and Ahrens*, 1971; *Arndt et al.*, 1971].

Gratz [1984], consistent with an earlier study [*Schall and Horz*, 1977], observed planar microfaults of thickness of ~ 1 to 10 microns oriented at 45 to 55 degrees to the shock wave propagation direction in approximate agreement with the orientation of the maximum resolved shear stress with respect to the shock propagation direction [*Jones and Graham*, 1971]. These microfaults are filled with glass; displacement across these microfaults is seen by offset of broken rutile needles separated by a microfault. The observations of steps emerging from the surface perpendicular to the shock direction in scanning electron micrographs [*Ashworth and Schneider*, 1985] are consistent with the production of microfaults characterized by relative displacements.

The microfaults are to be distinguished from shock lamellae [*Stöffler*, 1972,1974], the latter occur in sets of 4 to 12 and are narrow planes of lengths to 50 microns and thicknesses of 0.5 to 500 nm [*Gratz*, 1984]. Additionally, the shock lamellae have an orientation nearly perpendicular to the shock propagation direction and appear to be crystallographically controlled. Lamellae are apparently formed at the same time as the microfaults.

Alternatively, a series of tensile cracks, also filled with glass, and oriented perpendicular to the shock propagation direction are observed. These are features of the release of the material from high pressure. These

tensile remnants of the shock recovery experiment are smaller than the microfaults and the glass they contain appears to have originated in the microfaults and then migrated into the connected tensile cracks upon release (A. Gratz, personal communication).

Wackerle [1962] found that fused quartz shocked to near 50 GPa were recovered in single pieces. Furthermore, samples composed of several initially divided pieces of fused quartz were recovered as a continuous mass.

Ashworth and Schneider [1985] shock loaded quartz along the *a* axis to pressures of 30 GPa and although they make no mention of microfaults [*Anan'in et al.*, 1974a; *Gratz*, 1984] they note regions of patchy glass which could be microfaults when the differences in magnification of the studies are considered. Numerous glass filled lamellae of thicknesses less than 1 micrometer and a crystallographically controlled orientation, as seen by *Gratz* [1984], parallel to the shock direction were noted. The amorphous material within the lamellae is similar to that produced by the irradiation of quartz (metamictization).

The above observations are evidence for the proposed model of heterogeneous deformation as shown in Figure 2a after *Grady* [1980]. In the Figure, a shock wave is shown propagating towards the right in a brittle material. The shock compressed region on the left has experienced a heterogeneous compression and is comprised of blocks of solid material separated by layers of melt. The melt layers are oriented at 45 degrees to the direction of shock propagation (i.e., the direction of maximum shear stress). The solid blocks are subjected to volumetric compression due to application of a hydrostatic

stress via the regions of melt.

Figure 2b displays an enlarged view of a microfault. Most of the material within the confines of the microfault is melt of pressure dependent co-ordination. As cooling of the microfault occurs, high pressure phase begins to nucleate (black squares) within the region of melt (light grey region) similar to the observations of *Ashworth and Schneider* [1985]. In quartz, a change in co-ordination from 4 to 6 could be easily and rapidly be accomplished in a melt as suggested by the models of *Stolper and Ahrens* [1987]. Further, this nucleation of high pressure solid phase in equilibrium with the melt may constrain the temperature of the microfault to that of the fusion curve. The observed temperatures from mixed phase regions along the Hugoniot would be the melting or liquidus temperature of the equilibrium solid phase.

At a high enough shock pressure the volumetric density of these melted or sheared zones becomes large relative to the volume of the undamaged blocks such that rapid equilibration of the shocked material takes place on a time scale much less than the shock wave experiment. At this point the deformation would appear much more homogeneous [*Grady, 1977,1980*].

This model explains the loss of shear strength behind the shock front, and the phase transformation by shock wave in quartz.

EXPERIMENTAL

Experiments were conducted on the 40 mm powder gun at the Seismological Laboratory of the California Institute of Technology in a configuration similar to that of *Kondo and Ahrens* [1983a]. Planar shock waves were generated within the samples by impact of metal flyer plates (Table 1, Figure 3) with velocities of up to 2.5 km/s. Heating of the sample by shock compression resulted in the emission of visible light. This light propagates through the unshocked portion of the sample and is directed to a diffraction grating spectrometer whose output is recorded by a charge coupled device (EG&G Princeton Applied Research optical multichannel analyzer, model 1257) to give a visible spectrum. This digitally recorded spectrum is then used to determine a grey-body temperature for the material while it is under shock compression. In addition to the spectrometer, a photodiode supplies a record of intensity vs. time and a framing camera [*Schmitt et al.*, 1986] gives short exposure images of the sample under shock compression.

Cylindrical fused silica samples (T12 Optosil 1, Amersil Corp.) of 2.54 cm diameter and 0.6 cm thickness were purchased prepolished to an optical grade flatness. The specific gravity of the samples was measured to an accuracy of 0.02% by the Archimedean method. Silver was vacuum-deposited on one face to mask light flashes on arrival of the shock wave at the driver-sample interface due to gaps [*Kondo and Ahrens*, 1983b]. The silver-coated sample face was mounted with epoxy or phenyl salicylate to a metal driver plate (Table 1) which had been polished to a mirror finish with a final corundum grit of 0.3 micron. An opaque mask with an

hole of inside diameter of 13.0 mm was centered on the sample free surface to eliminate light emissions from regions near the edges of the sample.

Shock pressures were determined with the impedance match method [*Duvall and Fowles*, 1963] and using the Hugoniot data for fused quartz, magnesium, aluminum, copper, and tungsten in *Marsh* [1980] and the experimentally determined projectile velocities.

Light recorded by the optical multichannel analyzer passes through a series of photocathode intensifiers which are controlled by a high voltage pulse; the detector may thus be gated to record spectra on time scales as short as 100 ns. Gate times during an experiment are typically 300 to 500 ns and are normally timed to occur while the shock wave is near the center of the sample.

A portion of the beam is focused by a series of lenses into a TRW framing camera which images the sample during shock propagation. The framing camera allows up to three 100 to 500 ns exposures to be taken during propagation of the shock wave in the sample.

During setup of each experiment, the image of the fused silica sample is focused on the spectrometer slit (Figure 4). Magnification of the image was approximately 1:1, however this varied due to movement of the experiment chamber between successive experiments resulting in slightly different light paths which changed the focus. The image of the unmasked region of the sample was ~ 15 mm in diameter and covered the slit with a 13 mm height.

Calibration of the spectrometer was accomplished before each experiment with a standard of spectral radiance (Optronics model number

550, lamps IR-30 and IR-159). As small variations in current affect the radiance of the lamp considerably, current to the lamp was monitored to $\pm 0.03\%$.

Spectral radiances of the calibration lamp were supplied by the manufacturer (Figure 5). However, the spectrometer required a value of spectral radiance over each of the 500 channels representing a different wavelength. Appropriate values were supplied by a fit of the given calibration lamp radiances to the linearized Planck greybody radiation law (Equation 3). The curves through the known radiances represent the equivalent greybody function for the lamp, this function was used in the reduction of the spectrometer data. The maximum error introduced by use of the fits shown is below 2% over the most sensitive spectrometer wavelengths (550 to 750 nm).

During calibration, all controllable variables (light path, spectrometer delay (200 to 700 ns), and spectrometer gate (300 to 500 ns)) were identical to those of the experiment in order to reduce errors. A camera shutter set at 1/500 second exposure was used to trigger the system and to limit the amount of dark current into the spectrometer. Based upon tests using longer shutter open times it was found dark current radiance errors are less than 5% of the magnitude and independent of wavelength.

Light levels from the fused quartz samples during an experiment were less than 25% the level of light emitted by the calibration lamp. Since the CCD spectrometer had a limited light level sensitivity range, low light levels resulted in large signal to noise problems, whereas high light

levels could saturate the detector. Co-ordination of the calibrated light level with the expected light level for an experiment involved maximizing the amount of light arriving at the spectrometer during calibration with neutral density filters. The spectral characteristics of the neutral density filters were carefully accounted for by taking a number of tests of various light paths through the system.

The wavelength function of the spectrometer was determined by use of common spectral lines [Weast, 1982] of hydrogen (656.2 nm), helium (587.6 nm, 667.8 nm, 706.5 nm), oxygen (777.3 nm, 844.6 nm) and mercury (578.6 nm, 546.1 nm). Low intensities from the standard gas lamps made calibration at experimental gating times (~ 300 ns) difficult, several tests were conducted for each lamp to reduce the signal to noise ratio. The accuracy of the wavelength function changed with wavelength and was inversely proportional to the density of standard lamp lines used in a given region of the spectrum; the maximum uncertainty encountered in the positioning of the spectral lines was approximately 3 nm. Curvature of the vidicon intensifiers results in a variation of the wavelength interval per spectrometer channel, on average, the range was ~ 0.8 nm/channel.

A single photodiode was used to obtain a record of light intensity vs. time during the shock wave experiments. Output of this photodiode was simultaneously recorded with oscilloscopes and a LeCroy transient recorder which samples data every 10 ns. The photodiode and the spectrometer view different but adjacent areas of the sample (Figure 4).

DATA REDUCTION

The observed number of counts at each spectrometer channel is proportional to the time-integrated intensity of the emitted radiation at the given wavelength. To obtain a spectra, the raw data in counts were multiplied by a conversion factor relating spectral radiance intensity to number of counts derived from the calibration procedure (Figure 6). The temperatures of the standard lamps are approximately 2700 K, the calibration factor thus is very sensitive to errors towards the red and violet ends of the observed spectral range.

The determination of shock temperatures from the observed spectra assumes: 1) the material radiates as a Planck greybody, 2) the emissivity of the material under shock compression is independent of wavelength, 3) the temperature range of the experiment is such that a Wien's law spectral radiance distribution may be used. Assumptions 1) and 2) are based on the shape of the observed fused quartz spectra which, in general, are indicative of a thermal radiation: no line or band emissions are noted and the spectra monotonically decrease from the infrared to the ultraviolet. The complete Planck greybody function is given by:

$$R = \frac{\epsilon c_1}{\lambda^5 (\exp(c_2/\lambda T) - 1)} \quad (1)$$

where R is the spectral radiance, ϵ is the emittance, T is the temperature of the emitting body, λ is the given wavelength, and c_1 and c_2 are constants with values of $1.191 \cdot 10^{-16} \text{ Wm}^2/\text{steradian}$ and 0.01438 mK, respectively. For assumption (3) it is presumed that the simplification of Equation 1, given by the Wien's greybody function:

$$R = \frac{\epsilon c_1}{\lambda^5 \exp(c_2/\lambda T)} \quad (2)$$

is justified in order to facilitate linearization of the formula. Errors introduced by use of this formula are larger for higher temperatures and longer wavelengths; however, for the temperature range encountered in the experiments (2000 K to 4500 K) and for the range in wavelengths (450 nm to 880 nm) the maximum error is less than 2% and for most cases many orders of magnitude smaller. As a check on the validity of the linearized results, an iterative least squares routine was implemented with the full Planck greybody formula; results were always within 50 K of the linearized version.

The linearized Wien's law is given by:

$$\ln\left(\frac{c_1}{R \lambda^5}\right) = \frac{c_2}{\lambda} \frac{1}{T} - \ln(\epsilon). \quad (3)$$

where $\ln(c_1/(R \lambda^5))$ and c_2/λ are dependent and independent variables and a standard linear regression is then applied to find the slope $1/T$ and intercept $-\ln(\epsilon)$ which yield temperature and emittance, respectively.

Since data at either wavelength range extreme have a larger uncertainty due to the reasons mentioned above, each datum is relatively weighted by its corresponding raw data count (see Figure 6). The raw data function supplies a reasonable weighting scheme as it inherently accounts for both sensitivity of the system and spectral regions of maximum emission (i.e., lowest signal to noise ratio).

Formal standard deviations derived from the least squares routine yield temperature uncertainties on the order of ± 30 K. This is unrealistically small and on the basis of experience with the data

suggest an uncertainty of ± 200 K is more appropriate for the observed temperatures.

RESULTS

Figure 7 shows the spectra observed for x-cut crystalline shots #611 and at shock pressures of 10.6 and 7.6 GPa respectively. The shape of the spectra are greatly influenced by a non-thermal emission centered at approximately 500 nm and cannot be fit to a greybody curve. The shape of the spectra, however, is indicative of broadened line emission superimposed upon a greybody spectra. The determination of shock temperatures was the primary motivation for this study; further experiments on crystalline quartz were discontinued since in this pressure range the large non-thermal component of the x-cut quartz spectrum does not allow accurate temperature determination.

Brannon et al., [1983a] reported spectral and imaging data for shocked crystalline quartz from 6 to 26 GPa. They also observed a nonthermal bandlike spectral feature with peaks at 400 and 600 nm. Present observations displayed an emission at 500 nm, the difference in data may be due to the higher wavelength resolution of this study. The source of this nonthermal radiation is unclear and may be related to photoluminescence or cathodoluminescence [*Brannon et al.*, 1983b]. Since no nonthermal component was observed from fused quartz, the 500 nm emission of crystalline quartz may be related to breaking of long range order within the crystal or the piezoelectricity of quartz.

Another proposed explanation may be related to the production of the shock lamellae in crystalline quartz. Because fused quartz is a homogeneous, isotropic, amorphous material, it has no structural preferences and probably no shock lamellae are produced in fused quartz. Conceivably, thermal radiation may be associated with the shear heating within microfaults (in both fused and crystal quartz) and nonthermal radiation could arise from the formation of the diaplectic glass within the shock lamellae (in crystal quartz only).

The framing camera photograph consists of two exposures of shot #634 under shock compression (Figure 8), each of 100 ns duration and separated by 300 ns. The first exposure was initiated 50 ns after entry of the shock wave into the sample. The corresponding photodiode record for #634 (Figure 12) commences with zero intensity at the entry of the shock wave into the fused quartz sample and builds in a linear fashion with time. This furnishes convincing evidence that the bright spots observed in the photograph originated within the compressed fused quartz and not at the driver-sample interface.

Figures 9, 10, and 11 and Table 1 present the spectra and related data for the present experiments on fused quartz. These spectra are indicative of thermal radiation, no line or band emissions are observed as in crystalline quartz. The best fit Planck greybody function is plotted through each spectrum.

Figure 9 are the spectra obtained in the pressure range from 10 to 16 GPa which is the region of permanent densification along the fused quartz Hugoniot. Unexpectedly, the temperatures are highest in this lowest of

pressure ranges studied. Shots #688 and #634 display a marked drop in intensity at lower wavelengths (less than 450 nm). This is attributed to the poor sensitivity of the spectrometer in this wavelength region and further points out the importance of weighting the data in least squares fits of Equation 3. This effect may also be seen at the high wavelengths where there is a noticeable increase in the signal to noise ratio. The level of noise for shot #635 is particularly high as in this experiment most of the light was routed to the TRW framing camera.

Figure 10 displays spectra over the pressure range of 16 to 30 GPa, the spectra in this Figure are included on the basis of their emittance and their shock pressure within the mixed phase region of fused quartz. The two spectra in Figure 11 for shots #708 and #684 are within the same pressure range but both had emittances an order of magnitude larger and lower temperatures than the spectra of Figure 10.

Photodiode records for this series of experiments are shown in Figure 12. These are the combined results from both the oscilloscope and transient recorder data. Since the photodiode supplies only a qualitative response, all intensities have been normalized. The response of the photodiode is dependent upon the amount of light allowed to reach the spectrometer, the calibration lamp therefore gives a response of 200 mV to 700 mV and the typical peak amplitude during an experiment ranges from 100 mV to 300 mV. It should be noted that the photodiode and the spectrometer view different but adjacent regions of the sample during an experiment (Figure 4). The photodiode records are discussed below.

It should also be noted that shot #686 at 17.4 GPa had no observable light emitted during the accumulation time of the spectrometer. Additionally, the photodiode recorded essentially no intensity until 500 ns after entry of the shock wave into the sample. This is followed by a large flash when the shock wave reaches the free surface of the sample at approximately 1 microsecond.

Least square temperatures fit to Equation 3 for the above spectra are plotted in Figure 13 and the corresponding emittances vs pressure are plotted in Figure 14 and given in Table 1.

The range of the estimated Hugoniot temperatures of fused quartz is shown in Figure 13. The high estimate is that of *Kondo and Sawaoka*, [1981] and the low estimate is that of *Wackerle* [1962]. We believe these estimates to be extreme. It is difficult to calculate the continuum shock temperature of fused quartz under shock compression by the method of *Ahrens et al.*, [1969] because of the lack of knowledge of an accurate isentropic equation of state of fused quartz above 3 GPa and of the correct thermodynamic parameters. For example, fused quartz has an extremely low thermal Gruneisen parameter [*Stacey*, 1977] of 0.012 at 1 bar pressure due to its small coefficient of thermal expansion. This may not remain small with increasing pressure. Another complicating factor is the anomalous compression of fused quartz; the bulk modulus of fused quartz is known to decrease to approximately 2.3 GPa [*Kondo et al.*, 1981] and increases above this pressure [*Meade and Jeanloz*, 1987].

DISCUSSION

HETEROGENEOUS DEFORMATION IN FUSED QUARTZ

New evidence for the heterogeneous deformation of fused quartz under shock compression is supplied by the framing camera photograph of shot #634 (Figure 8) in which is seen a pattern of bright regions or hot spots in agreement with earlier studies [Brooks, 1965; Brannon *et al.*, 1983b]. The number of regions grows with time, i.e., with the distance the shock wave has propagated into the fused quartz. Additionally, these regions persist and grow for at least 500 ns: the time spanning this set of exposures. Further, that these localized hot zones originate in the fused quartz is proven by the simultaneously recorded photodiode record (Figure 12). No sample-driver interface flash is seen; light intensity commences at zero when the shock wave enters the fused quartz and grows linearly with time suggesting a constant production of shear bands with time as the shock propagates into the sample.

INTERPRETATION OF TEMPERATURES AND EMITTANCES

Figure 13a displays the observed shock temperatures vs. pressure. The estimated fields of silica melt (L), stishovite (St), coesite (Co), crystalline quartz (Q), and cristobalite (Cr) are shown. The stability fields below 5 GPa are from Jackson [1976] phase boundaries are from Davies [1972] and Jeanloz and Thompson [1983]. Temperature measurements above 30 GPa of Sugiura *et al.*, [1982] are also included as well as the calculated fused quartz shock continuum temperatures

[*Kondo and Sawaoka*, 1981; *Wackerle*, 1962] as discussed above.

The following observations are made from this plot:

- 1) Temperature appears to decrease with pressure from over 4400 K at 10 GPa to approximately 3000 K at 30 GPa. This decrease levels off above 16 GPa.
- 2) Observed color temperatures exceed those calculated for the continuum by nearly 1000 K.
- 3) three anomalous shots above 25 GPa exhibit lower temperatures.

There is an approximate dependence on emittance on pressure (Figure 14). Below 25 GPa all emittances are less than 0.02. A threshold pressure appears to be reached above 25 GPa where the emittance increases by at least an order of magnitude. Finally, two shots have emittances which are almost those of a blackbody. The emittance may be an approximate measure of the areal density of the localized zones of deformation; the density of the features of heterogeneous deformation increases with pressure [*Ashworth and Schneider*, 1985]. The approach of the emittance towards unity at the higher pressures may signal the onset of a regime of a more homogeneous shock deformation reached when the density of microfaults becomes high.

The observation of temperatures greatly in excess of the estimated continuum temperatures and corresponding low emittances suggest that fused quartz undergoes a process of heterogeneous shock deformation as also suggested by similar shock wave experiments of crystalline halite, calcite, and gypsum [*Kondo and Ahrens*, 1983a], and potassium

chloride, lithium fluoride, and periclase [Schmitt *et al.*, 1986; Schmitt and Ahrens, 1983].

Based upon observations of the observed temperatures and emittances the data has been divided between three fields, each of which is characterized by the type of shock deformation occurring, and the magnitude of the emittance.

The first field is within the pressure range of 10 to 16 GPa which has been referred to as the permanent densification region of fused quartz [Sugiyura *et al.*, 1981]. The spectra within this field have low emittances (less than 0.015) and the highest temperatures (3200 K to 4400 K) which decrease with increasing pressure. This anomalous behavior is not understood but appears to be related to the transformation to stishovite which commences near 7.5 GPa at low temperatures as noted earlier. The high temperatures may be characteristic of the transformation at the initiation of the mixed phase region. This phenomenon has been more fully observed in crystalline NaCl over the B1-B2 solid phase transition region [Schmitt *et al.*, 1987]. However, a direct analogy between NaCl and SiO_2 may not be justified since the phase relations of NaCl are much simpler than those of SiO_2 which has a number of high pressure solid phases.

On a speculative note, these high temperatures may be related to the large activation energy or energy barrier ($\sim 3.5 \cdot 10^6 \text{ J/kg}$ of SiO_2) to the reconstructive quartz to stishovite transformation [Grady, 1980; Holm *et al.*, 1967] which dictates the amount of energy required to initiate the transformation. Deposition of insufficient energy by the shock wave to activate the transformation must raise the temperature of the material.

The maximum temperature rise ΔT may be estimated by the amount of energy required to activate the process A divided by the heat capacity C_p of the material or:

$$\Delta T = \frac{A}{C_p}. \quad (4)$$

Using the values of activation energy and the heat capacity of Table 2 the maximum temperature rise before activation of the transformation is estimated to be approximately 4700 K; this figure is consistent with the observed radiative temperatures of ~ 4400 K near 10 GPa. If the observed emittances (~ 0.01) near 10 GPa yield the relative volume of material undergoing the phase transformation, then the process mentioned above requires approximately 10 to 15% of the total energy available in shock compressed fused quartz by the Rankine-Hugoniot energy equation:

$$\Delta E = \frac{1}{2}(V_o - V_1)P_h. \quad (5)$$

where ΔE is the internal energy change upon shock compression, V_o and V_1 are the initial and final specific volumes, respectively, and P_h is the shock pressure.

Field #2 contains the shots which have low emittances (less than 0.1) and are within the mixed phase region of the fused quartz Hugoniot. The high temperatures and low emittances indicate a heterogeneous deposition of thermal energy within the fused quartz. These temperatures would represent the melting temperature of stishovite at high pressure. The melting temperature of stishovite between 20 and 30 GPa is thus estimated to be $3000 \text{ K} \pm 200 \text{ K}$. This melting temperature is

independent of pressure to the degree of precision available in this experiment.

Lyzena et al., [1983] measured shock temperatures in both fused and crystalline quartz between 60 and 140 GPa. At these pressures silica is believed to be well into the stishovite phase (see earlier discussion of phase transformation in quartz). Based on their temperature measurements, Lyzena and co-workers provided evidence for the melting of stishovite and obtained an hypothetical melting curve at high pressure (Figure 13b). The present data lie on or close to the proposed melting line of *Lyzena et al.*, [1983] thus supporting the present hypothesis that the melting temperatures are observed.

Field #3 contains the two high emittance (near blackbody) - lower temperature data indicative of the onset of a more homogeneous deformation. These temperatures are close to those observed by *Sugiura et al.*, [1982] by a similar technique above 30 GPa. Higher emittances appear to indicate more homogeneous shock deformation according to *Lyzena and Ahrens* [1980] who report emittances > 0.75 and in most cases near unity for fused quartz shocked to pressures in excess of 70 GPa. *Lyzena et al.*, [1983] further demonstrate the homogeneity of the deformation at these pressures in a streak camera record of alpha quartz at 106.5 GPa which displays uniform emittance.

The reason for the overlap of the pressure ranges of fields #2 and #3 is not known and may be related to differences in sample preparation or due to the inherent difference in fused quartz samples. Within a given sample of fused quartz the variation of the index of refraction, which is a measure of density, can be as high as $5.0 \cdot 10^{-6}$ (*Heracus-Amersil*

Specifications). Also, small bubbles and inclusions do exist within the fused quartz; differences in their density may result in the observed shot to shot variations.

PHOTODIODE RECORDS

Ideally, the present experiments are designed to prevent light generation at the driver-fused quartz interface. Because of the opaque silver film, detected light originates only within the shocked fused quartz. The production of hot shear deformation zones appears to be nearly constant with time. The hot zones persist for a timescale on the order of the transit time of the material [Grady, 1980], and the absorption characteristics of the shock compressed material appear to be nearly identical to the as yet unshocked material. This results in a linear increase in the emitted intensity with time [Boslough, 1985]; the ideal photodiode record would therefore have a nearly linear increase in intensity with the propagation of the shock wave into the sample. Such behavior is displayed by shots #634 and #709 at pressures of 9.8 and 27.5 GPa, respectively.

However, many of the experimental photodiode records (Figure 12) display a more complicated structure which includes evidence of some fused quartz sample-driver interface heating (coinciding with entry of the shock wave into the sample). This problem has been addressed by Urtiew and Grover [1974], Grover and Urtiew [1974], Boslough [1985], and Svendsen *et al.*, [1986], who demonstrate this is a potential problem with measurements of shock temperatures in clear or translucent materials. Briefly, their model involves deposition of a quantity of heat at the interface

due to collapse of micron scale gaps between the driver and sample. This heat results in the production of high temperatures at the interface which decay inversely with $(\kappa t)^{1/2}$, where κ is the thermal diffusivity of the metal driver which is assumed to have a higher thermal conductivity with respect to the sample. The observed photodiode records display this behavior. Note that the rise time of the photodiode due to its amplification is approximately 50 ns. Thus because of the finite rise time of the photodiode electronics, the peak signal is nearly coincident with the entry of the shock wave into the sample.

Some photodiode signals are thus a superposition of the interface gap flash and the light produced within the fused quartz [Boslough *et al.*, 1986] (Figure 15). In Table 1, severe contamination occurs when spectrometer accumulation occurs primarily during the time of decreasing photodiode intensity due to interface cooling. Moderate contamination occurs when the gap flash is low amplitude and part of the accumulation time overlaps with decreasing photodiode intensity. Contamination is considered to be small when accumulation time occurs in an increasing or almost level portion of the light intensity signal observed after the interface temperature has decayed to the continuum temperature of the shocked sample. In Table 1, none refers to no gap flash contamination.

Although we expected that the potential effect of gap flashes on the measured temperatures might be very deleterious, we found surprisingly that shots at nearly the same pressure yield similar temperatures regardless of the degree of gap flash contamination. For example, shot #634 at 9.8 GPa (with no gap flash) and shot #688 at 10.0 GPa (severe gap

flash) gave temperatures in excess of 4000 K and similar emittances. Additionally, shots #710 and #711 at pressure of 29.6 and 30.2 GPa with severe and small gap flashes, respectively, both gave temperatures near 2900 K and emittances between 0.2 and 0.3. A buffering of the gap temperatures to the melting point of the compressed silica (via mechanisms similar to those proposed within the microfaults may be occurring). Another explanation may be related to the evolution of the shock wave with increasing penetration into the sample. *Hayes and Graham* [1978] calculate stress-time histories for an impact of quartz on quartz and assume dependent shear stress relaxation occurs. The energy density dissipated at a given depth within the sample decreases monotonically until a stable shock wave is formed. Higher temperatures and densities of deformation features occur near the interface.

MICROFAULT TEMPERATURE VARIATION WITH TIME

The observed temperatures, which we here suggest to be at the temperature of the melt in the microfaults at shock pressure in excess of 20 GPa, could be dependent on the scale of the deformation features on account of heat conduction. *Heinz and Jeanloz* [1987] noted the effect of spatial variations in temperature upon measured radiative temperatures in melting experiments within high pressure diamond cells. We now examine the conditions where the observed radiative temperatures will not represent the melting temperature in the microfault.

We use the theory for solidification of a dike emplaced in a body of cooler country rock [*Turcotte and Schubert*, 1982]. Consider

the initial microfault created within the shock wave front zone as a layer of thickness d of molten silica at the melting temperature for a given pressure. Initially, material within the shear band is assumed to be totally molten and at the melting temperature T_m . Material outside of the shear band is shock compressed solid at the cooler temperature T_o which is nearly equal to the continuum temperature. In most cases $T_m \gg T_o$. Cooling of the shear band and the release of the latent heat of fusion during freezing occurs due to heat flow from the microfault zone at T_m to the compressed solid at T_o by phonon or radiative conduction.

The temperature, T , in the solid is given as a function of y , the distance away from the interface and time, t as:

$$\frac{(T - T_o)}{(T_m - T_o)} = \frac{\text{erfc}(\nu)}{\text{erfc}(-\lambda)} \quad (6)$$

where

$$\nu = \frac{y}{(2\kappa t)^{1/2}} \quad (7)$$

and the constant λ for a given set of thermal parameters must be determined by numerical calculation in the transcendental equation:

$$\frac{L(\pi)^{1/2}}{(C_p(T_m - T_o))} = \frac{e^{-\lambda^2}}{\lambda(1 + \text{erf } \lambda)} \quad (8)$$

T is the temperature as a function of time t and position y from the microfault. The liquid-solid shear band boundary is at $y = 0$ for $t = 0$. κ is the thermal diffusivity which is a function of the heat capacity of the material C_p , the coefficient of thermal expansion: α , and the density of the material. L is the latent heat of fusion of silica, and erf and

$erfc$ are the error function and $1 - erf$, respectively.

The above equations have been solved using thermodynamic parameters for fused silica (Table 2) and for differences between the melt in the shear band and the surrounding solid of 500 K, 1000 K and 2000 K. Figure 16a illustrates the growth of a frozen layer of solid with time for varied temperature differences. The Figure illustrates that for an extreme temperature difference of 2000 K the solidified zone thickness is approximately 1.6 microns 1 microsecond after creation of the microfault. Figure 16b displays the temperatures profile with time for an initial temperature difference of 1000 K. Radiation emitted from these cooler regions will have an effect on the spectra of the observed light. The magnitude of this effect is dependent on the ratio of the volume of the cooled temperature material to the volume of the initial melted material. At maximum, the timescale for these experiments is 1 microsecond. Figures 16a and 16b therefore suggest that for a microfault of thickness of 10 microns, which is expected on the basis of shock recovery work on quartz [Gratz, 1984], this cooling by conduction and freezing will have little effect upon the observed temperatures. However, if the thickness of the microfault is 5 microns or less the cooling will have a much greater effect on the observed spectra. A microfault of 5 or less microns created early on in the shock experiment may freeze solid by the end of the experiment (Figure 17). The effect of this cooling on measured temperatures, however, is offset by the constant production of microfaults with time as the shock wave propagates through the sample. Spectra are obtained from light accumulated not more than 800 ns after entry of the

shock wave into the sample.

The above modelling of microfault cooling should not be considered as the true history of a microfault. Shock recovery observations favor the nucleation of small particles of solid stishovite as depicted in Figure 2b [Ashworth and Schneider, 1985]. Additionally, the present model does not take into account the kinetics of crystallization. It appears that only small amounts of a silicate can nucleate from the melt on the time scale of a shock experiment [Rigden *et al.*, 1987]. Quartz crystallization from a melt as a function of pressure is not yet well understood; however, recent experimental work on the crystallization of quartz from fused quartz suggests that crystallization rate increases with pressure up to 4 GPa [Fratello *et al.*, 1980].

COMPARISON OF NEW MELTING DATA WITH EARLIER RESULTS

The proposed estimates (Figure 13a, Table 1) for the melting temperature of silica between 20 and 30 GPa (stippled area of Figure 18) are compared to melting data for Mg_2SiO_4 to 15 GPa [Ohtani and Kumazawa, 1981], $MgSiO_3$ to 4.6 GPa [Boyd *et al.*, 1964], $(Mg_{0.9}Fe_{0.1})SiO_3$ from 18 to 60 GPa [Heinz and Jeanloz, 1987], and SiO_2 to 2.5 GPa [Jackson, 1976]. According to Figure 18, both SiO_2 and $(Mg_{0.9}Fe_{0.1})_2SiO_4$, as well as the extrapolation of the melting data for Mg_2SiO_4 , are nearly the same temperature.

Using the melting temperatures of Figure 18 and phase diagrams at the pressures of 1 atmosphere and 2.5 GPa the solidus temperatures in the system Mg_2SiO_4 - SiO_2 are estimated. The room pressure phase diagram (Figure 19a) after Chen and Presnall [1975] and Bowen and Anderson [1914]

displays a peritectic point at a composition of 60.6 weight % SiO_2 . With increasing pressure, the composition of MgSiO_3 becomes an endmember component with the appearance of a Mg_2SiO_4 - MgSiO_3 eutectic (Figure 19b) [Chen and Presnall, 1975]. Based on studies of two phase systems [Kennedy and Higgins, 1972; and Newton *et al.*, 1962], the depth of the eutectic trough is expected to increase with increasing pressure, i.e., the minimum difference between the melting temperatures of a pure endmember component and the eutectic composition is observed at low pressure. In the extrapolations to higher pressures, the depth of the eutectic troughs are assumed to be the same as in Figure 19b, thus the eutectic temperatures at higher pressures represent a maximum value. Additionally, it is assumed in the extrapolations that the composition of the two eutectics of Figure 19b is constant, this assumption appears to be valid for the MgSiO_3 - SiO_2 eutectic [Chen and Presnall, 1975] but may be in error by $\sim 3\%$ for the other eutectic as the hydrous Mg_2SiO_4 - MgSiO_3 system has a eutectic composition near 53 weight % SiO_2 [Kato and Kumazawa, 1986]. Using the depth of the eutectic well from the MgSiO_3 melting temperature as determined by Chen and Presnall [1975] at 2.5 GPa (Figure 19b) and extrapolating to higher pressures with the melting temperatures of Figure 18, we infer the maximum solidus temperatures in the Mg_2SiO_4 - SiO_2 system would be ~ 2500 K at 12.5 GPa (Figure 19c) and 2650 K at 20.0 GPa (Figure 19d). Inclusion of the component FeO would further depress the solidus temperatures; for an estimated mantle Fe/(Fe + Mg) ratio of 10% [Jeanloz and Thompson, 1983], the solidus temperature may be further depressed by ~ 130 K as estimated from the drop in the solidus temperature in the system Mg_2SiO_4 - Fe_2SiO_4 at one

atmosphere [Bowen and Schairer, 1935]. The maximum solidus temperatures within the mantle may thus be ~ 2370 K at 12.5 GPa and 2520 K at 20.0 GPa.

Weertman [1970] suggests that the effective viscosity of a solid at a constant creep rate is given by:

$$\eta = C \exp \left(g \frac{T_m}{T} \right) \quad (9)$$

where η is the effective viscosity, C is a constant, T_m is the melting temperature, T is the temperature of the material, and g is a dimensionless constant empirically found to be ~ 18.0 for most metals. The applicability of Equation 9 to brittle minerals is unknown; recent work on polycrystalline olivine [Green and Borch, 1987], however, indicates that at constant homologous temperature T/T_m the flow stress is independent of pressure to 2.5 GPa. The ratio of the proposed melting curve of SiO_2 (T_m) (Figure 18) to an upper mantle geotherm (T) (Figure 18) [Turcotte and Schubert, 1982] at pressures between 6.5 and 20 GPa is nearly constant near 0.6 suggesting little variation in the effective viscosity from 200 to 600 km within the Earth in agreement with the inferences of Heinz and Jeanloz [1987].

CONCLUSIONS

The major findings of this study are:

- 1) Fused quartz, like other brittle minerals subject shock compression, undergoes heterogeneous shock deformation characterized by localized zones of melting between 16 and 30 GPa. Evidence for this is supplied by the framing camera photographs that show hot regions which persist

and grow during shock compression and the high temperature-low emittance greybody spectra observed between 16 and 30 GPa.

2) The observed temperatures reflect the temperature of the localized zones of deformation. Shock recovery and release wave studies suggest that these zones are molten during shock compression. Shock recovery work and the Hugoniot of fused quartz suggest that the high pressure solid phase (stishovite) is produced within these shear bands or microfaults.

3) The melt temperature within these microfaults may be buffered by the production of the high pressure phase and as such may reflect the liquidus temperature of stishovite at high pressure. This suggests that this temperature is approximately 3000 K in the pressure range of 20 to 30 GPa and is independent of pressure.

4) Shock compression becomes more homogeneous with increasing pressure as evidenced by the general increase in emittance with pressure.

5) Shock temperatures observed in the pressure range of 10 to 16 GPa are up to 2000 K higher than temperatures at increased shock pressure and appears to be related to the onset of the fused quartz-stishovite phase transformation. These high temperatures and subsequent decrease in temperature with increasing pressure are not understood. Similar behavior is seen in NaCl over the B1- B2 phase transition region.

6) Solidus temperatures for Mg_2SiO_4 - SiO_2 system inferred from the present proposed melting curve for SiO_2 at high pressure together with other melting data suggest that the maximum solidus temperatures within the mantle are close to 2370 K at 12.5 GPa and 2520 K at 20.0 GPa. Comparison of the proposed SiO_2 melting curve with reasonable predicted geotherm

temperatures suggests that the effective viscosity between 200 and 600 km within the mantle is nearly constant.

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Table 1: Experimental Parameters

| Shot # | Flyer | Driver | Flyer Vel. (km/s) | Pressure (GPa) | Temp. (°K) | Emittance | I.F.E. | Field |
|--------|-----------------|--------|---------------------------|-------------------|---------------|---------------|----------|-------|
| 634 | Al _a | Al | 1.44 ± 0.02 | 9.8±0.2 | 4390±500 | 0.012±0.004 | none | 1 |
| 635 | W | W | 1.543 ± 0.02 | 15.2±0.3 | 3590±200 | 0.0047±0.0005 | small | 1 |
| 684 | W | Mg | 2.471 ± 0.05 | 25.7±0.7 | 2190±200 | 0.056±0.006 | small | 3 |
| 685 | W | Mg | 1.995 ± 0.03 | 20.0±0.4 | 2950±200 | 0.0058±0.0006 | severe? | 2 |
| 686 | W | Mg | 1.750 ± 0.02 | 17.4±0.4 | ^b | - | - | - |
| 687 | W | Mg | 1.400 ± 0.02 | 13.7±0.3 | 3240±200 | 0.0025±0.0003 | moderate | 1 |
| 688 | Al | Mg | 1.488 ± 0.02 | 10.0±0.2 | 4210±200 | 0.0064±0.0006 | severe | 1 |
| 689 | W | Mg | 2.492 ± 0.05 | 25.8±0.6 | 2930±200 | 0.055±0.006 | small | 2 |
| 691 | W | Mg | 2.295 ^c ± 0.04 | 22.2±0.6 | 3270±200 | 0.0035±0.0004 | small | 2 |
| 695 | W | Mg | 2.086 ± 0.03 | 21.0±0.5 | 2620±200 | 0.0023±0.0002 | small | 2 |
| 708 | W | Al | 2.455 ± 0.05 | 29.3±0.7 | 2030±200 | 0.780±0.08 | moderate | 3 |
| 709 | W | Al | 2.346 ± 0.04 | 27.5±0.6 | 2400±200 | 0.041±0.004 | none | 2 |
| 710 | W | Cu | 2.492 ± 0.05 | 29.6±1.0 | 2930±200 | 0.030±0.003 | severe | 2 |
| 711 | W | Al | 2.507 ± 0.04 | 30.2±0.5 | 2883±200 | 0.022±0.002 | small | 2 |

I.F.E. : Interface Flash Effect

^a : W = Tungsten, Al = Aluminum, Cu = Copper, Mg = Magnesium.

^b : No detected light emitted

Table 2: Thermal Parameters used in Calculations

| Symbol | Definition | Value | Reference |
|----------|--|--|--------------|
| C_p | Heat Capacity | 753 J/°Kkg | Weast (1982) |
| ρ | Density | 2204 kg/m ³ | Measured |
| L | Latent Heat of Fusion | 2.37×10^5 J/kg | Weast (1982) |
| k | Thermal Conductivity | 1.38 W/m °K | Weast (1982) |
| κ | Thermal Diffusivity($k/\rho C_p$) | 8.33×10^{-7} m ² s ⁻¹ | Calculated |
| A | Activation Energy of α - quartz to stishovite reconstructive transformation | 3.5×10^6 J/kg | Grady (1980) |

FIGURE CAPTIONS

Fig. 1a. Shock pressure vs volume data for crystalline (filled symbols) and fused (open symbols) quartz. Static compression curves for alpha quartz (Qz) (initial density = 2.65 g/cm^3 , bulk modulus = 37.1 GPa , bulk modulus pressure derivative = 6.0) and stishovite (St) (initial density = 4.288 g/cm^3 , bulk modulus = 305 GPa , bulk modulus pressure derivative = 4.5) [Kuskov and Fabrichnaya, 1987] calculated with Birch-Murnaghan finite strain theory are shown for comparison. The fused quartz and crystalline quartz Hugoniot coincide above 30 GPa suggesting that they are in the same phase state. Key to regions along Hugoniot: E - elastic deformation below HEL, PD - permanent densification of fused quartz, MP - mixed phase, St - stishovite, T - stishovite to melt transition under homogeneous shock melting, HM - homogeneous shock melting.

Fig. 1b. Shock wave velocity vs particle velocity for crystalline and fused quartz. Symbols have same meaning as in Fig. 1a.. Fused quartz data from Lyzenga et al., [1983] - (open triangles), Jackson and Ahrens [1979] - (open squares), Wackerle [1962] - (open pentagons), and Marsh [1980] - (open circles). Crystal quartz data from Lyzenga et al., [1983] - (filled 5 point star), Ahrens and Rosenberg [1968] - (filled inverted triangle), Grady et al., [1974] - (filled circle),

Fowles [1967] - (filled 4 point star), *Marsh* [1980] - (filled pentagon), *Trunin et al.*, [1971] - (filled square), and *Wackerle* [1962] - (filled hexagon).

Fig. 2. Schematic of model of heterogeneous shock compression in brittle materials, after *Grady* [1975, 1980]. **a.** Large scale view with shock wave magnitude P_h propagating towards the right in a sample. Behind the shock wave material is compressed and heterogeneous yielding has occurred on planes at 45 degrees to the direction of shock propagation. **b.** Microfault containing melt is surrounded by undamaged material. Upon cooling high pressure phase may nucleate.

Fig. 3. Shock radiation experimental configuration.

Fig. 4. View of spectrometer slit. Images of fused quartz sample and filament of calibration ribbon lamp shown. Light accumulated by spectrometer during experiment originates only within the region of the sample which is focused directly over the slit. Region of sample to right of slit is reflected by a mirror to the photodiode.

Fig. 5. Spectral radiances of calibration lamps. Triangles: lamp #1 (2680 K), circles: lamp #2 (2713 K).

Fig. 6. a. Raw data: response of spectrometer during experiment. **b.** Spectrometer response to calibration lamp. **c.** Output of spectral radiance lamp. **d.** Observed spectra derived from **a.**, **b.**, and **c.**

Fig. 7. Spectra observed for shots #611 and #615 for x-cut crystalline quartz. A large non-thermal component to the spectra

centered near 500 nm makes temperature determination infeasible.

Fig. 8. Framing camera photograph for fused quartz shot #634 at 9.8 GPa. Two exposures each of 100 ns duration and separated by 300 ns are shown.

Fig. 9. Observed spectra from fused quartz shocked between 9.8 and 16 GPa (Field #1 in Fig. 13a. and Fig. 14.).

Fig. 10. Observed spectra from fused quartz shocked in the mixed phase region (20 to 30 GPa) (Field #2 in Fig. 13a. and Fig. 14.).

Fig. 11. Observed spectra from fused quartz undergoing homogeneous shock deformation based upon low temperatures and high emittances (Field #3 in Fig. 13a. and Fig. 14.).

Fig. 12. Normalized intensity vs time, fused quartz. Fiducials a, b, and c represent times of initiation of spectrometer light accumulation, end of accumulation, and first shock wave arrival at sample free surface, respectively.

Fig. 13a. Fused quartz temperatures vs pressure. Data is separated into three fields: 1, 2, and 3, as in Fig. 14. Coesite-Stishovite-liquid triple point based on calculations of *Davies* [1972] and *Jackson* [1976]. Quartz-Coesite line based on data of *Holm et al.*, [1967]. Coesite-Stishovite phase line from *Jeanloz and Thompson* [1983]. Estimated shock temperatures of *Kondo and Sawaoka*, [1981] (K) and *Wackerle* [1962] (W). Shock temperature measurements of *Sugiura et al.*, [1982] are filled squares. Estimated stability fields of silica melt (L), stishovite (St), coesite (Co), crystalline quartz (Q), and

cristobalite (Cr) are shown.

Fig. 13b. Temperature vs pressure plot to include the data of *Lyzenga et al.*, [1983] and their proposed melting line. Plot contains only data from field #2.

Fig. 14. Log(greybody emittance) vs pressure. Field numbers are the same as in Fig. 13a..

Fig. 15. a. Hypothetical photodiode record assuming light originates only at driver-sample interface due to heating. **b.** Photodiode record assuming light originates only from features of heterogeneous deformation within the sample, **c.** Photodiode record obtained by superposition of records **a.** and **b.** **c.** is similar to those observed in the experiments.

Fig. 16. a. Solidified zone thickness vs. time after formation of microfault. $\Delta T = T_m - T_o$.

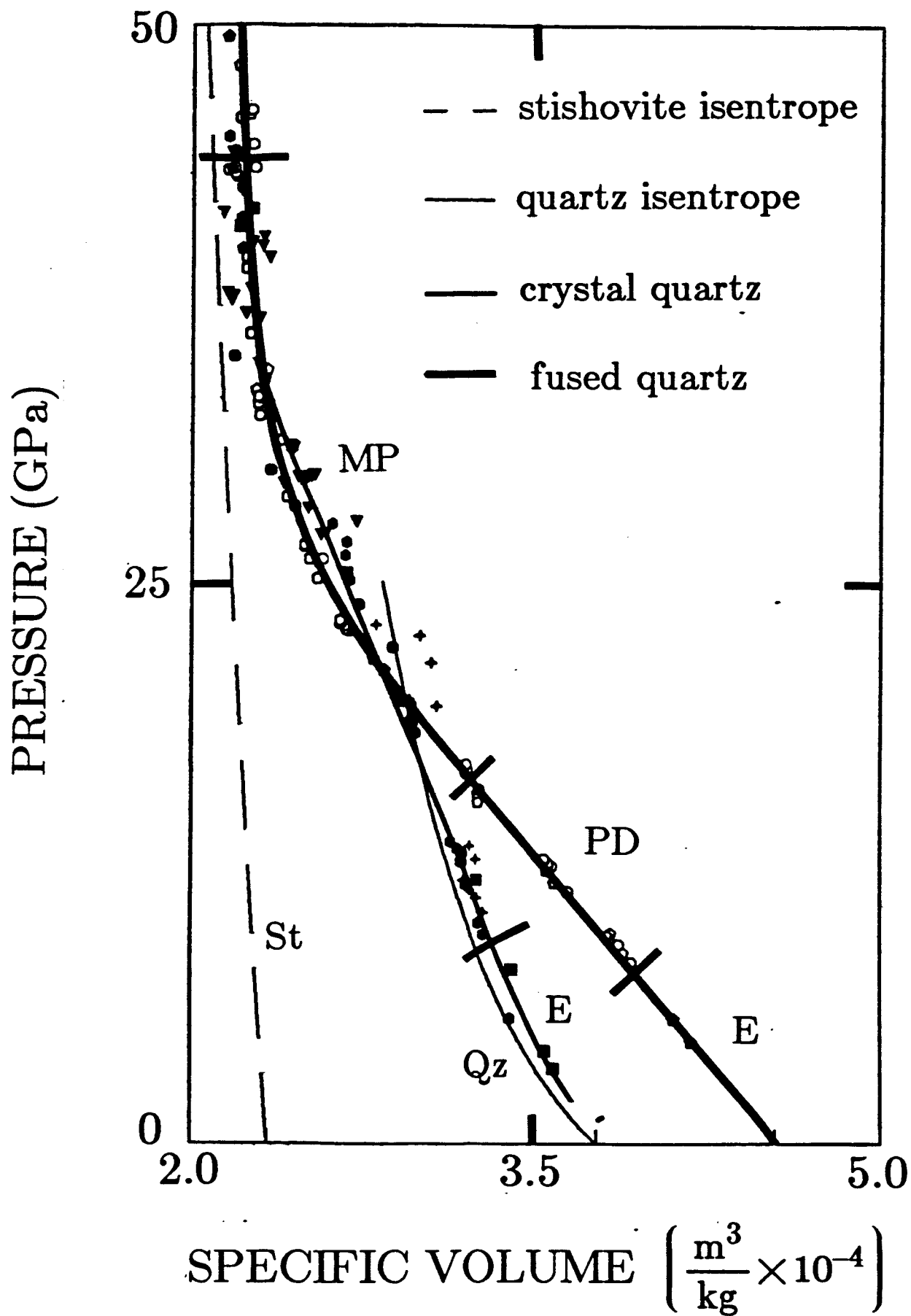
Fig. 16. b. Temperature distribution in the region of the initial microfault boundary at various times.

Fig. 17. Evolution and cooling history of a microfault for various times. At $t=0$, microfault is a layer of melt (shaded grey) at T_m surrounded by solid compressed material (hatched area) at T_o , while at t_2 and t_3 partial and complete crystallization (areas filled with close vertical lines), respectively, is indicated.

Fig. 18. Comparison of proposed SiO_2 fusion curve (stippled region) with measured and inferred melting data for SiO_2 to 2.5 GPa [*Jackson*, 1976], MgSiO_3 to 5.0 GPa [*Boyd et al.*, 1964], Mg_2SiO_4 to 15

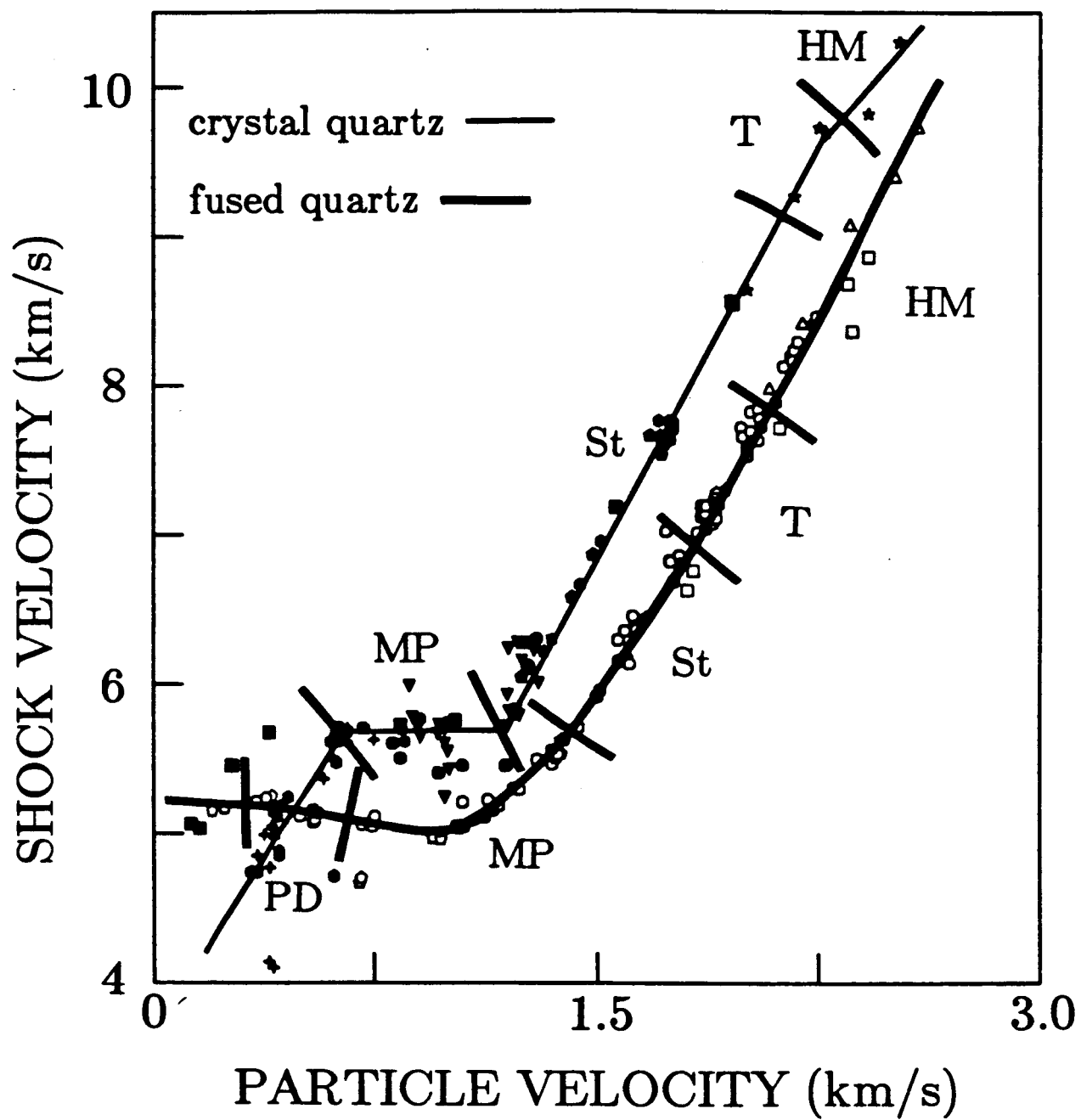
GPa [*Ohtani and Kumazawa*, 1981], and $(\text{Mg}_{0.9}\text{Fe}_{0.1})\text{SiO}_3$ from 18 to 38 GPa [*Heinz and Jeanloz*, 1987]. Inferred lower mantle adiabats from *Jeanloz and Richter*, [1979] and upper mantle adiabats from *Turcotte and Schubert*, [1982].

Fig. 19. a. Measured $\text{Mg}_2\text{SiO}_4\text{-SiO}_2$ phase diagram at 1 atmosphere after *Chen and Presnall* [1975]. L represents melt region. **b.** $\text{Mg}_2\text{SiO}_4\text{-SiO}_2$ phase diagram at 2.5 GPa. **c.** Proposed $\text{Mg}_2\text{SiO}_4\text{-SiO}_2$ phase diagram at 12.5 GPa. Depth of eutectics as measured from MgSiO_3 melting temperature are the same as in **b.** **d.** Proposed $\text{Mg}_2\text{SiO}_4\text{-SiO}_2$ phase diagram at 20 GPa. Depth of eutectics as measured from MgSiO_3 melting temperature are the same as in **b.**



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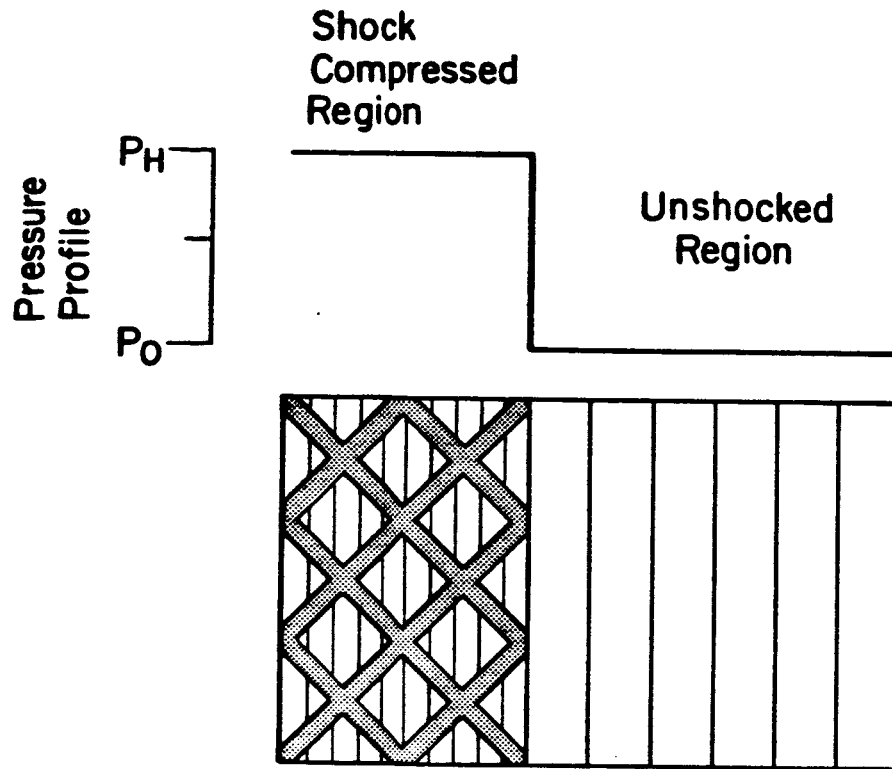
Figure 1a



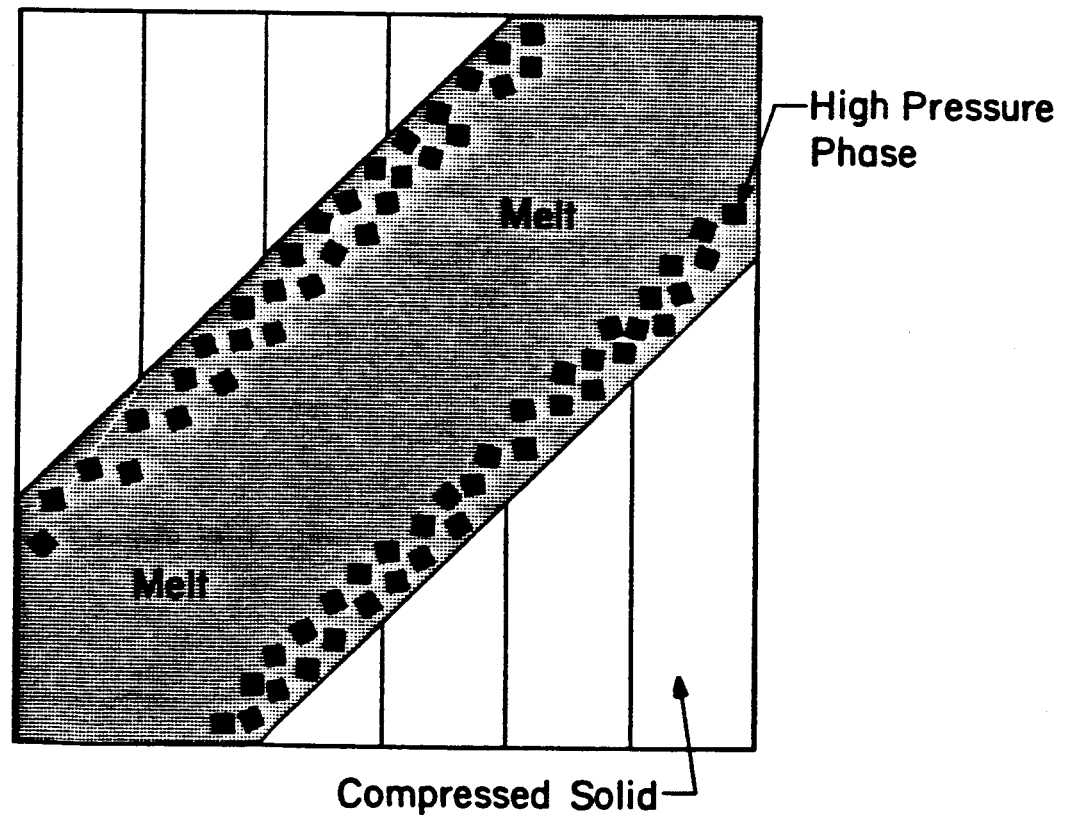
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Figure 1b

a)



b)



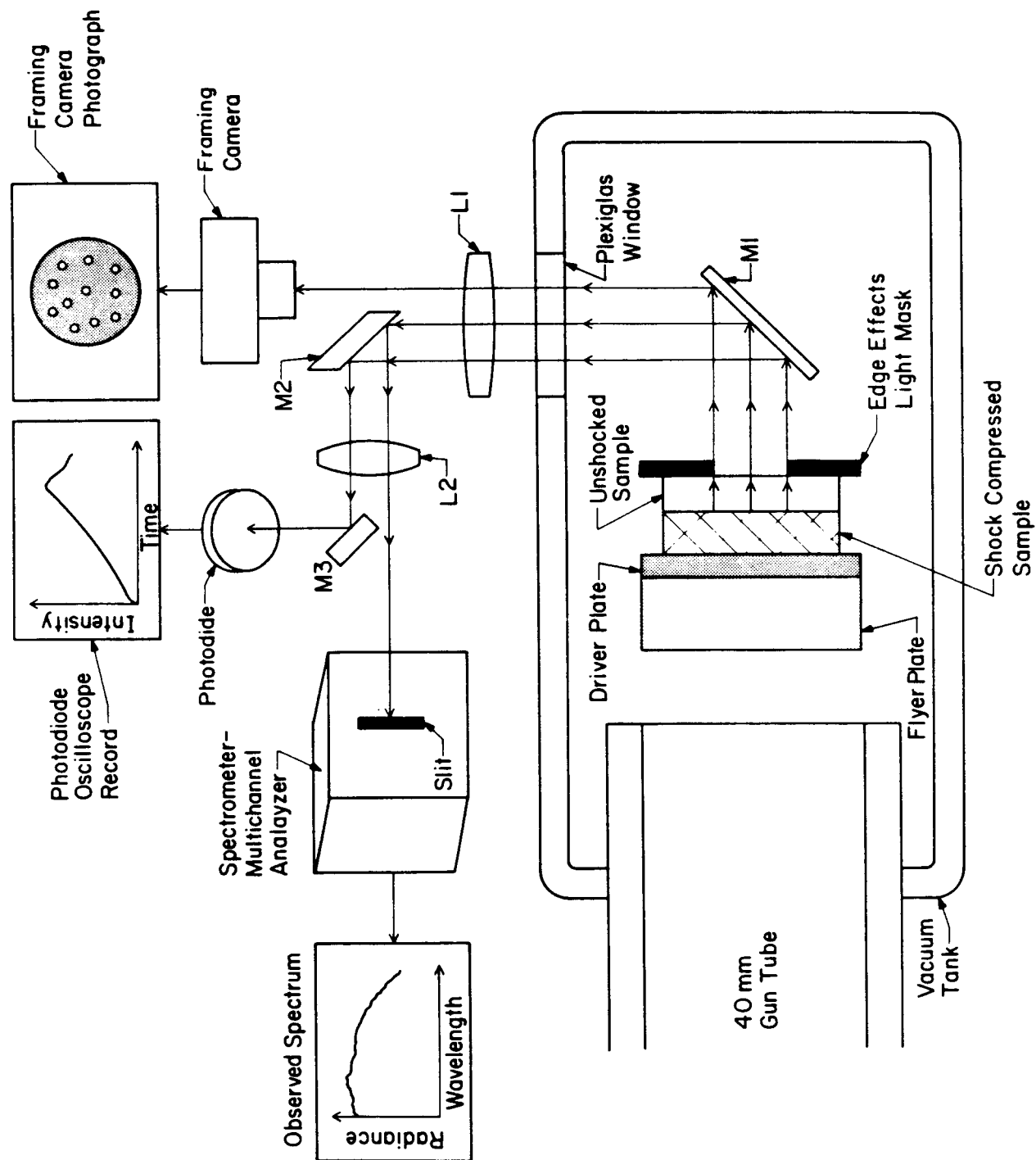
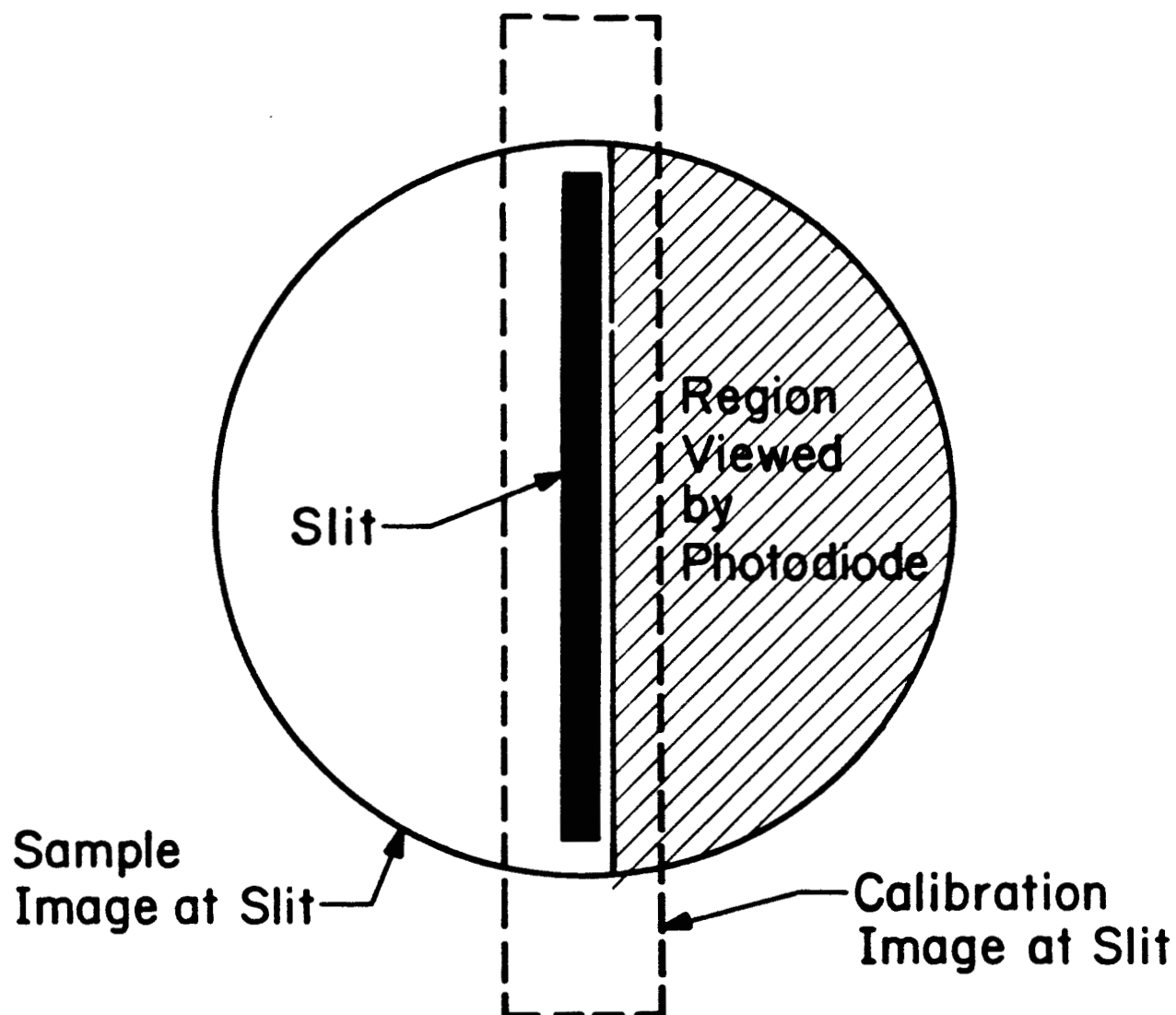
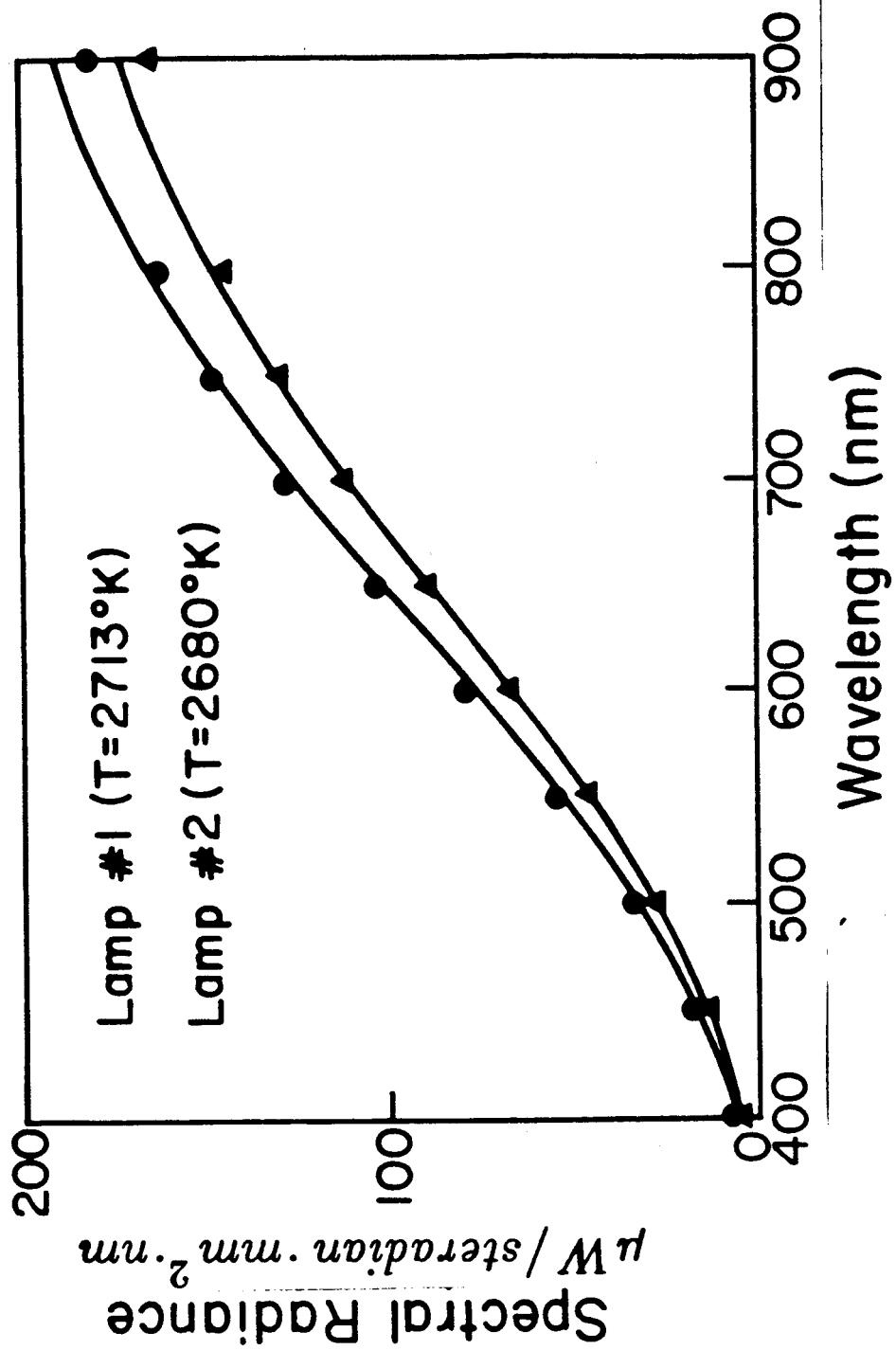


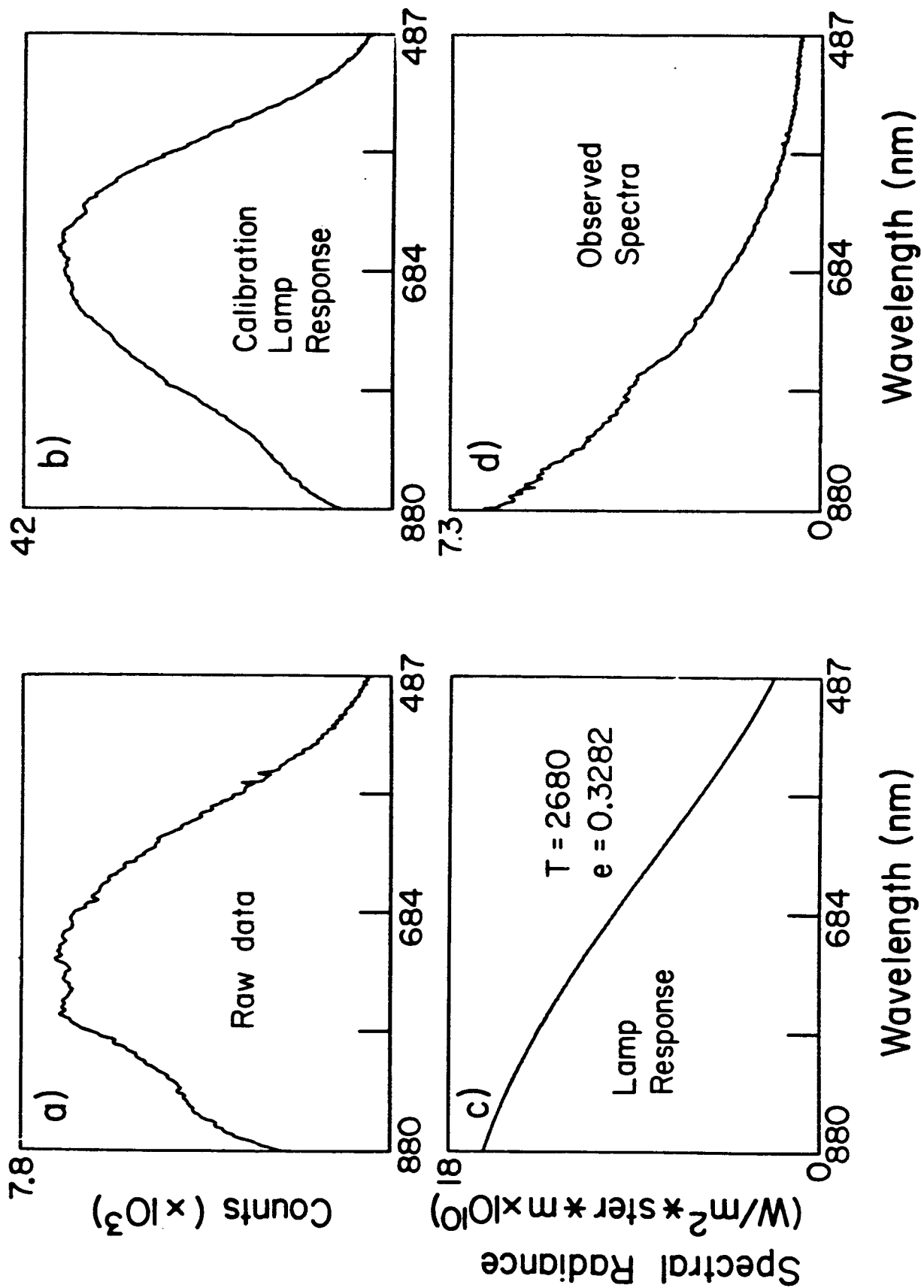
Figure 3

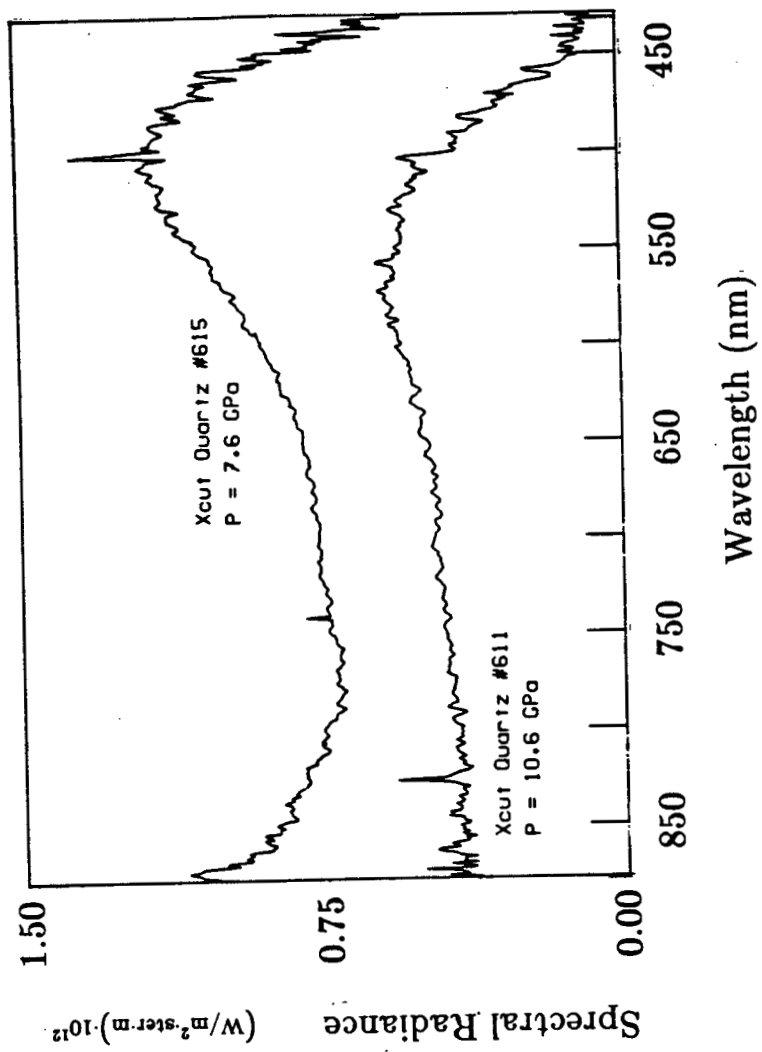


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Figure 4







FUSED SiO_2

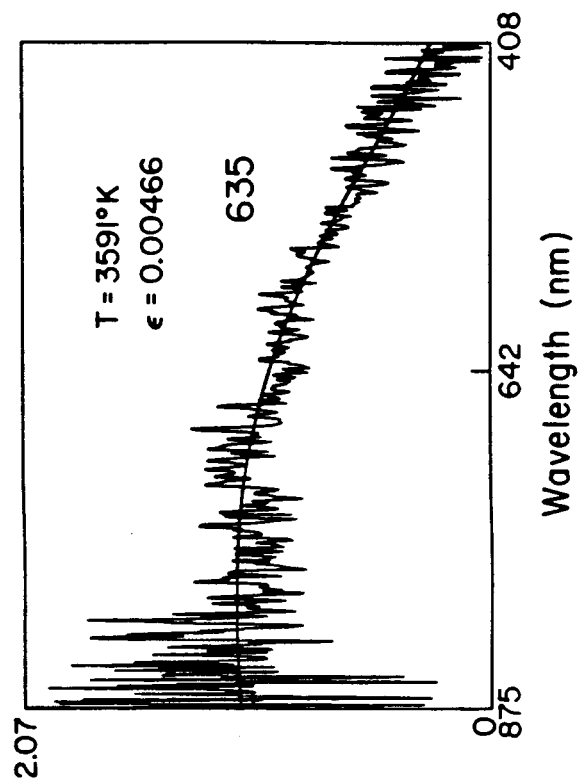
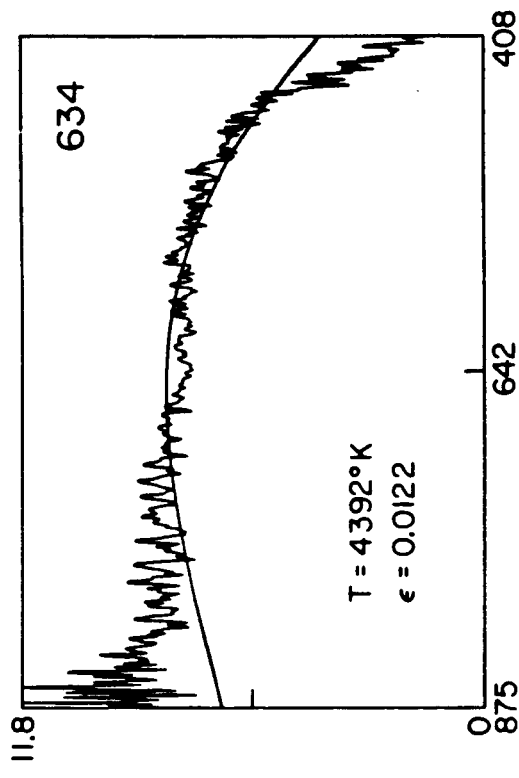
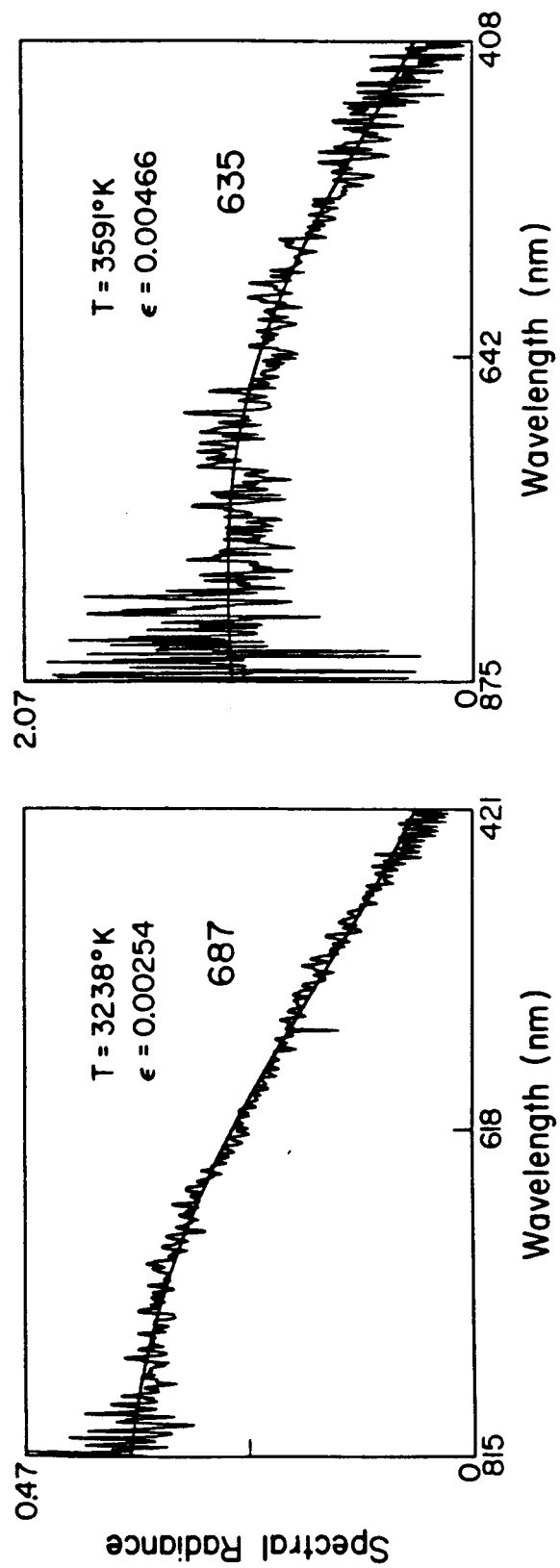
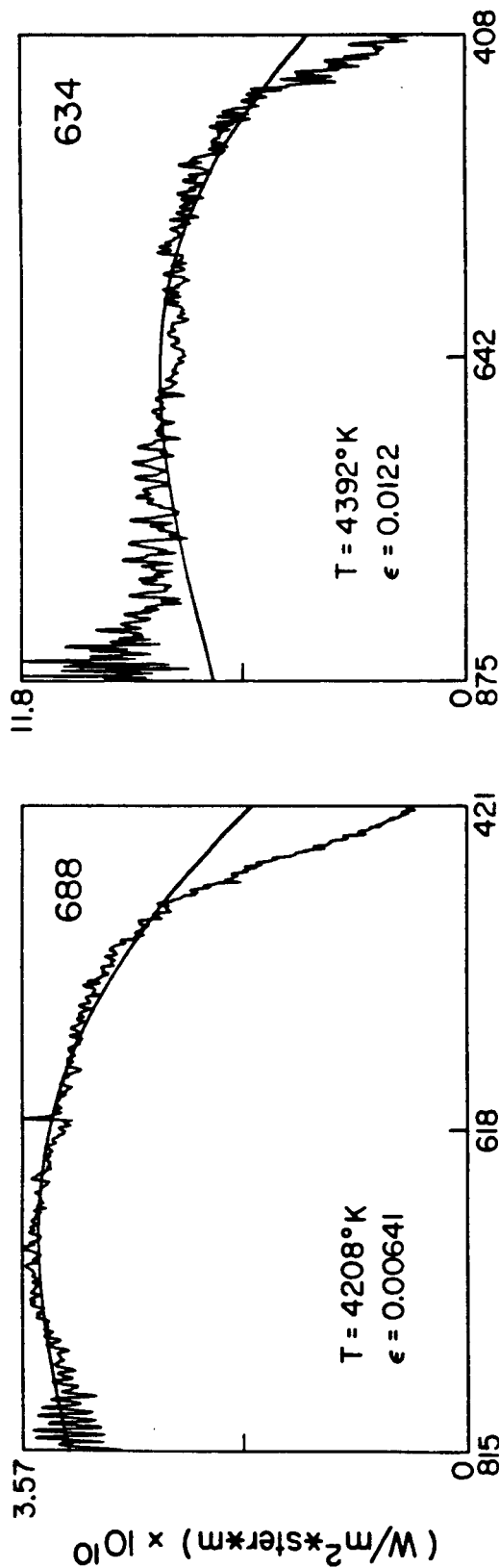


1.3 cm

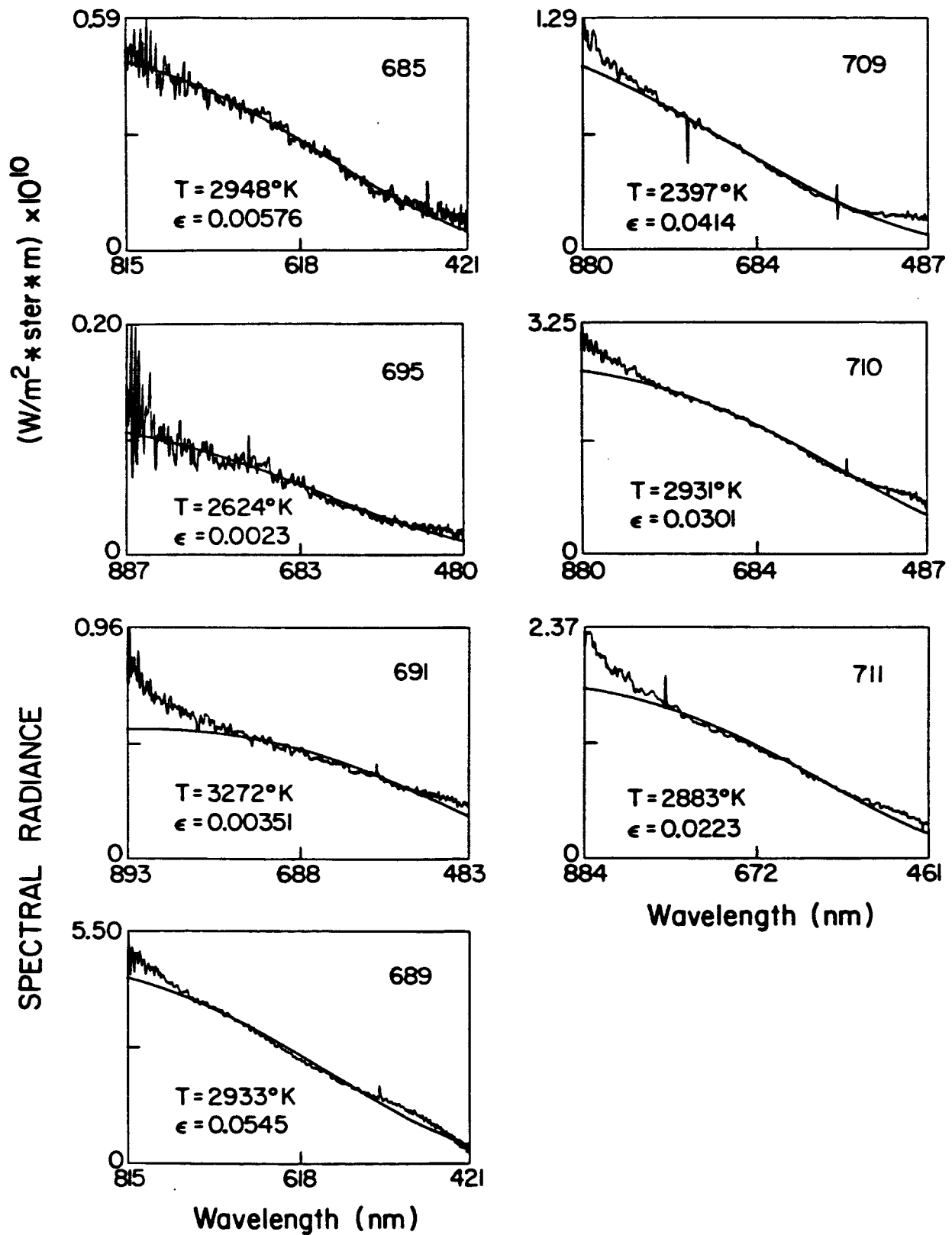
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Figure 8

FUSED QUARTZ PERMANENT DENSIFICATION REGION



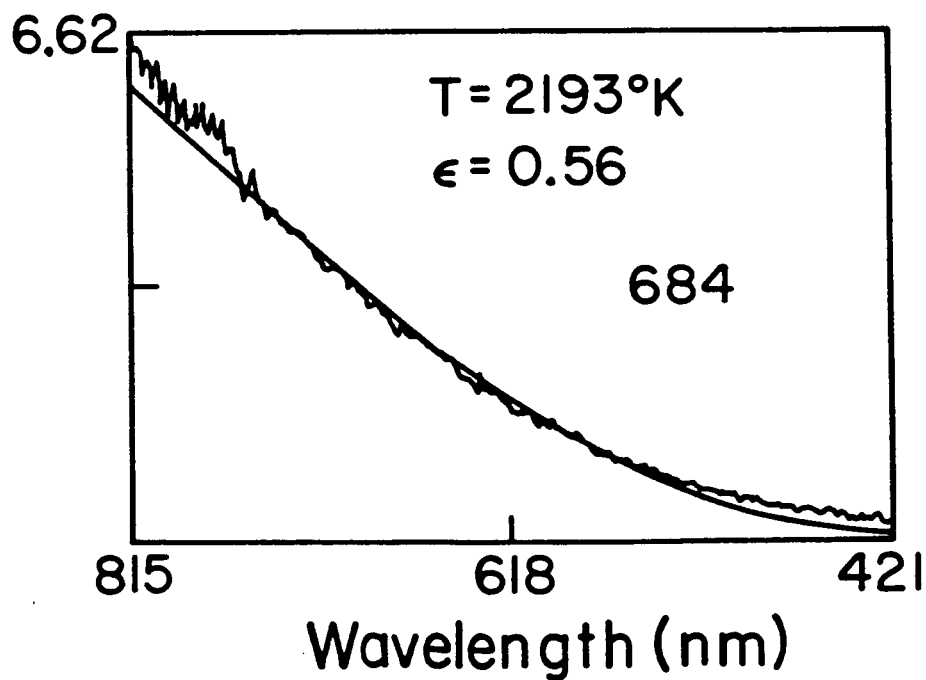
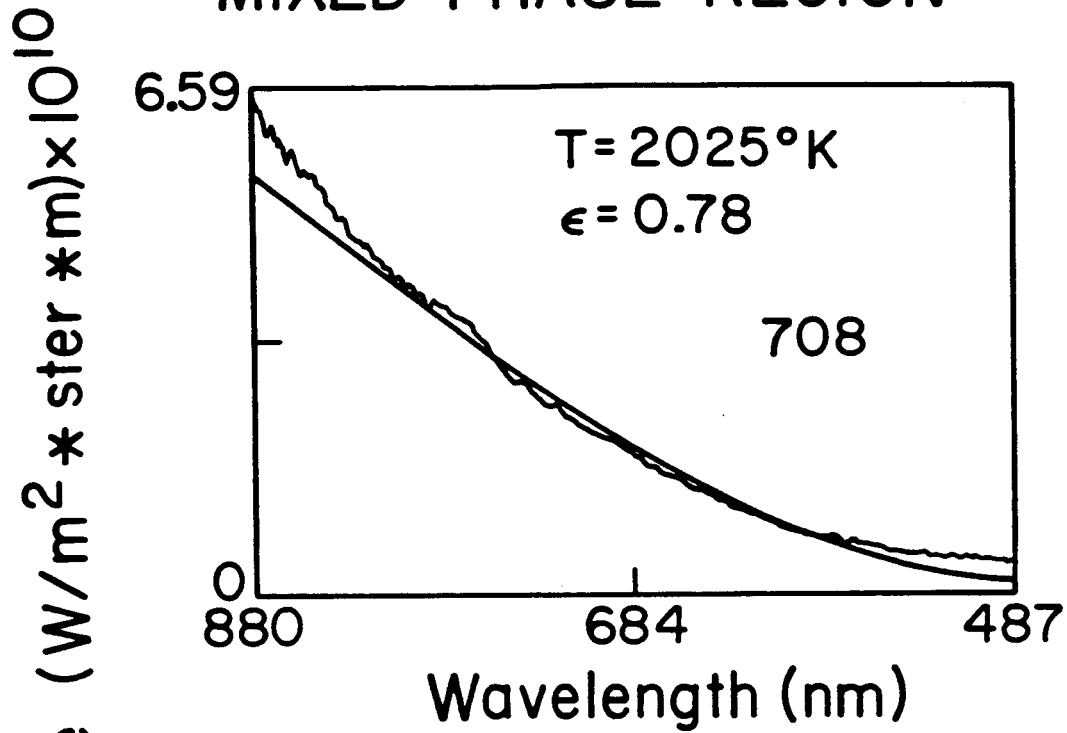
FUSED QUARTZ MIXED PHASE REGION

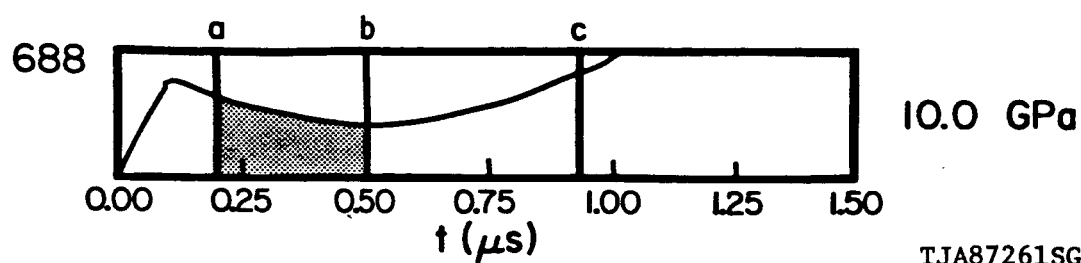
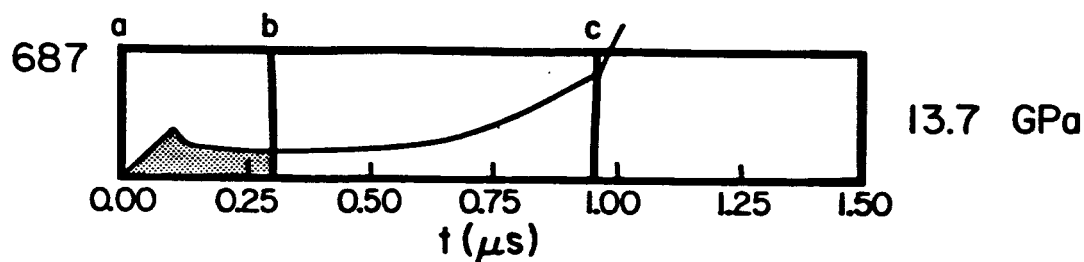
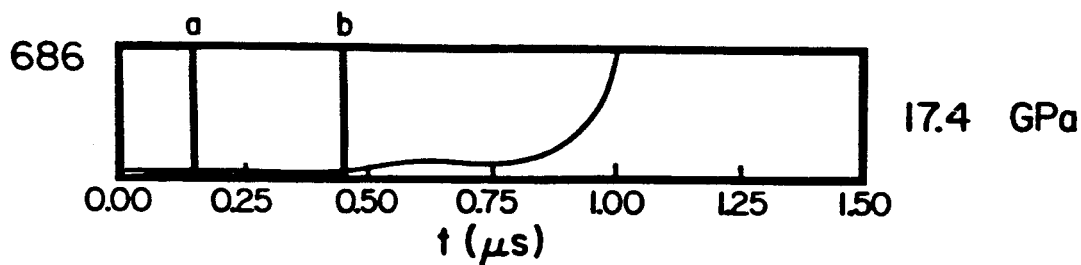
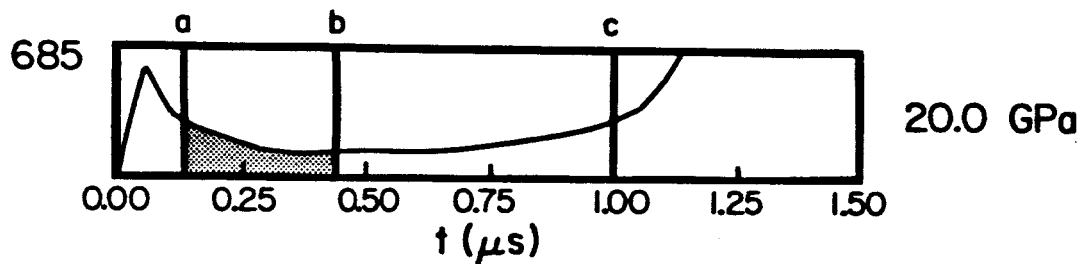
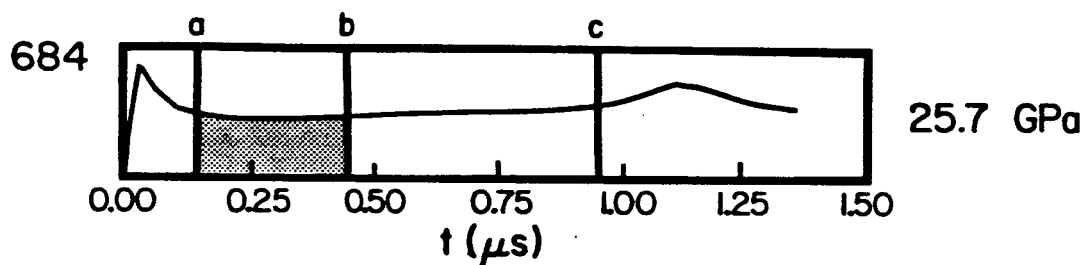
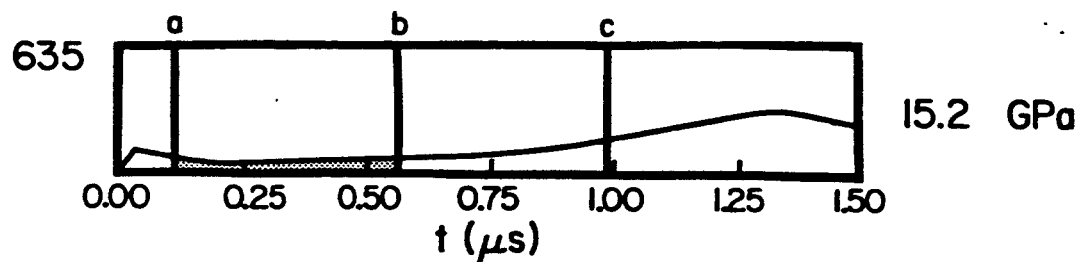
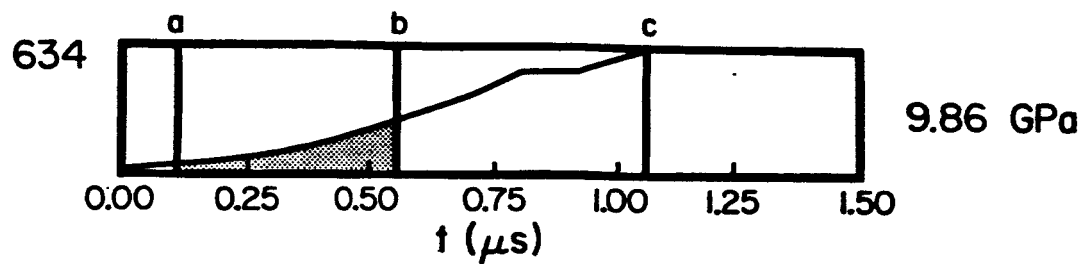


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Figure 10

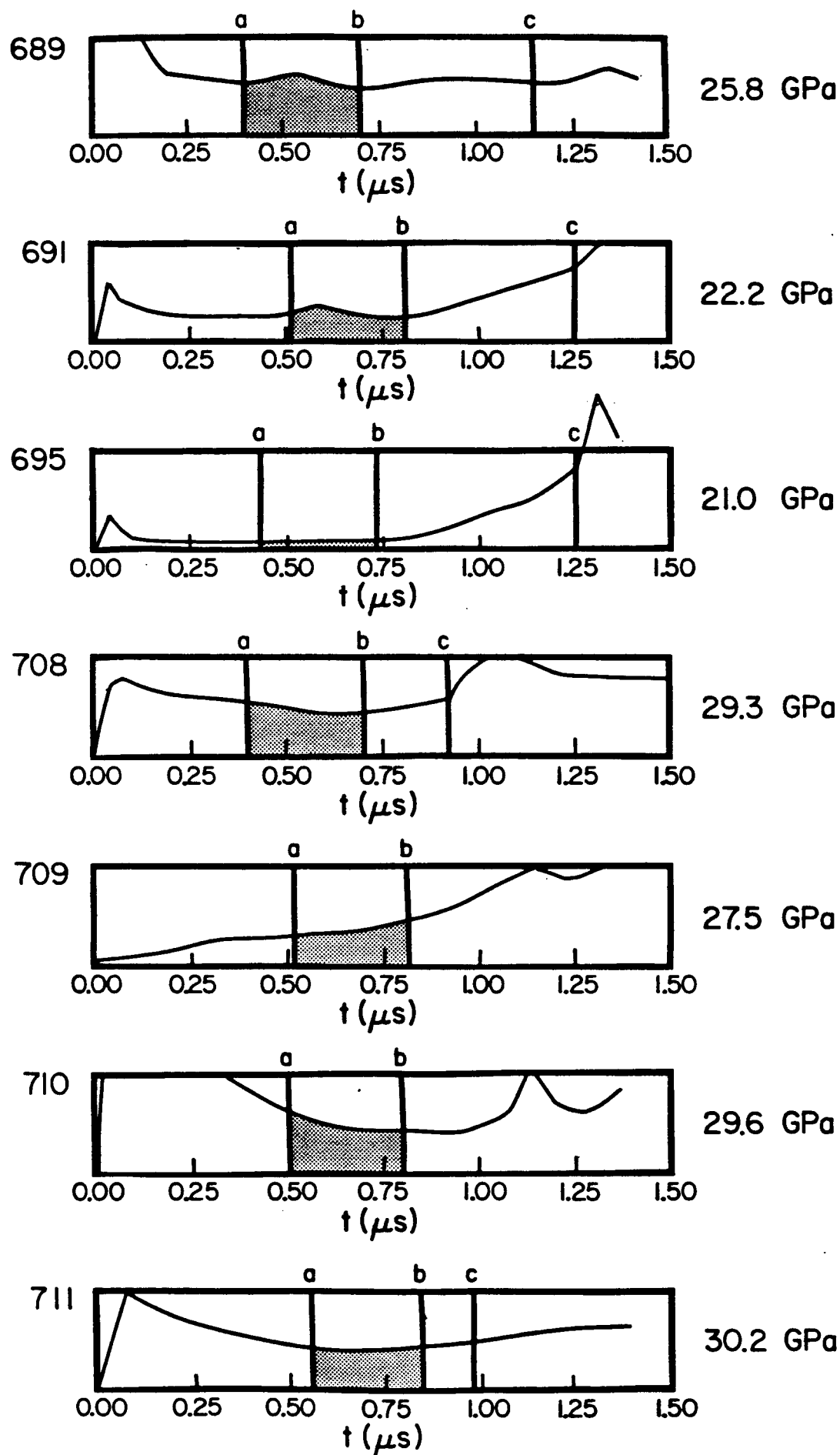
FUSED QUARTZ MIXED PHASE REGION





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Figure 12 1/2



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LFD

Figure 12 ²/₂

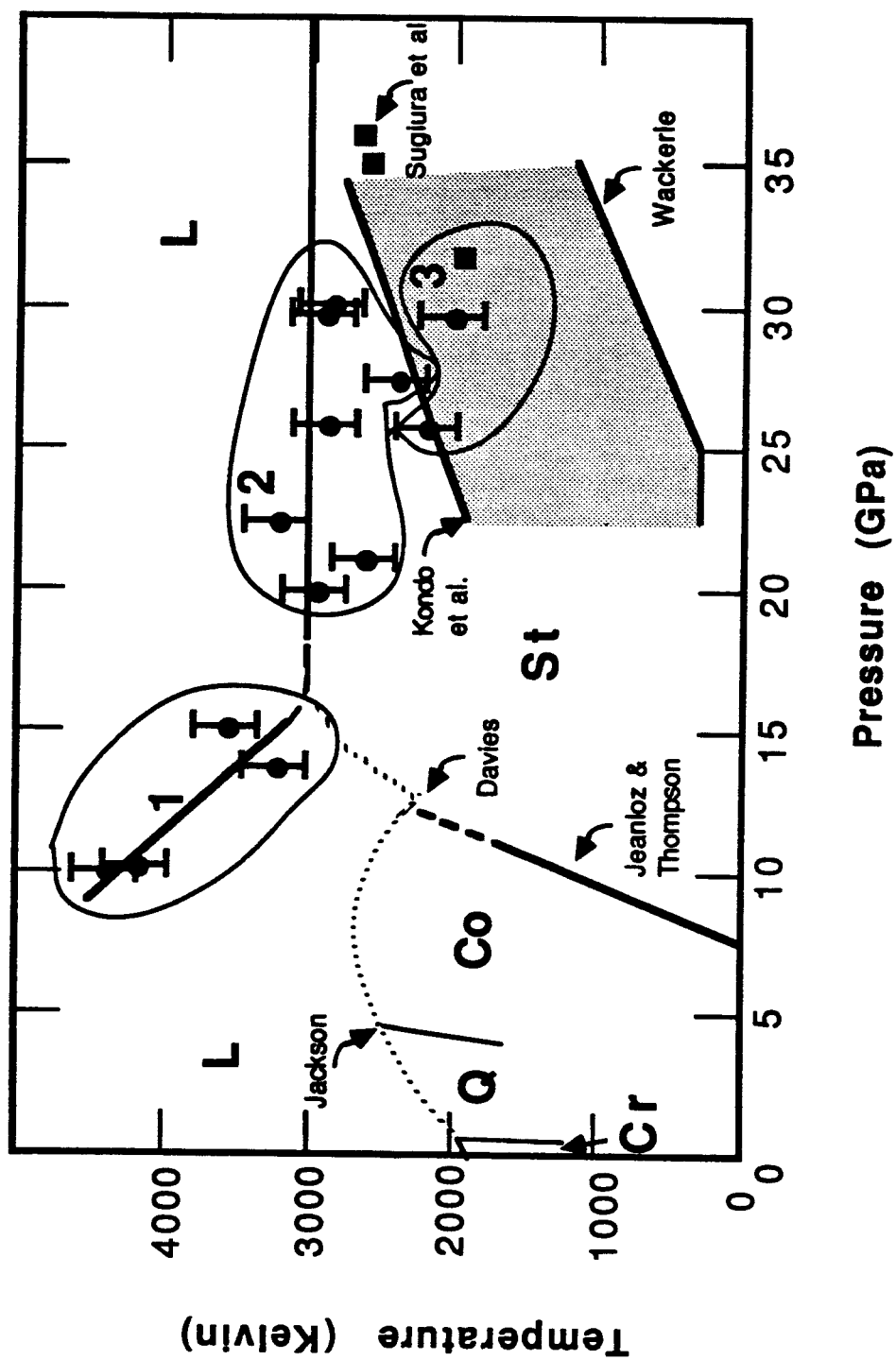


Figure 13a
Schmitt & Ahrens, 1987
to JGR

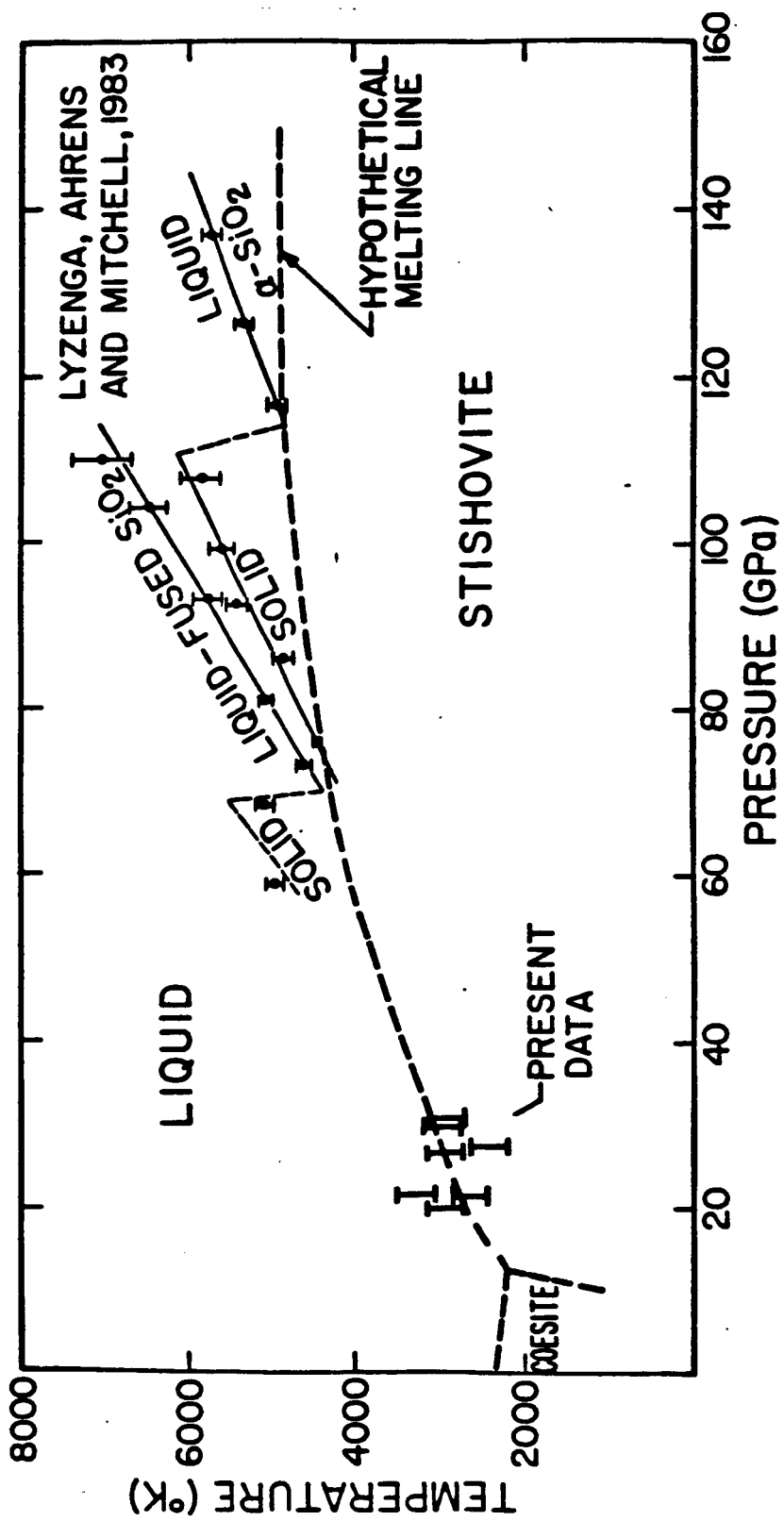
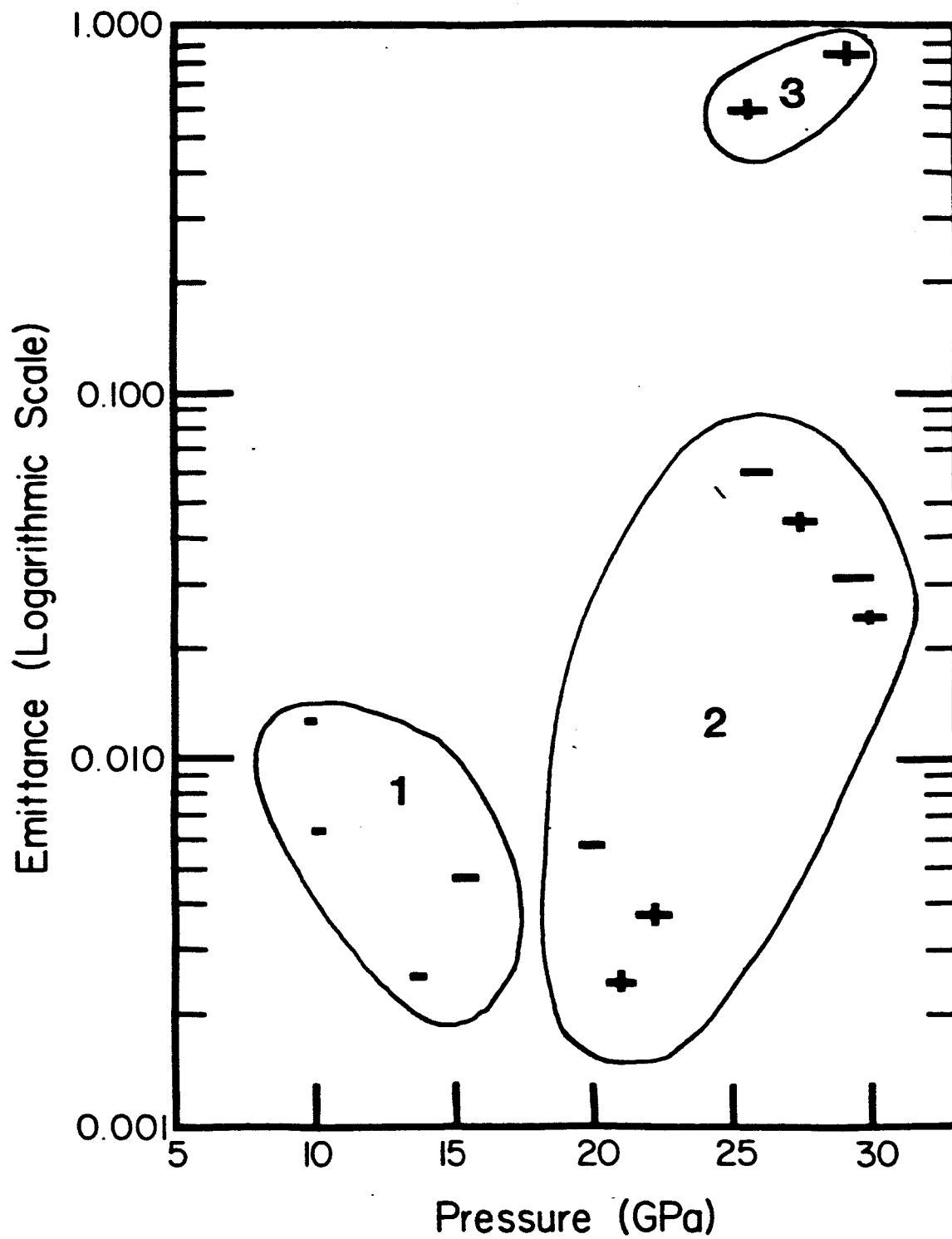


figure 13b

FUSED QUARTZ EMITTANCE VS. PRESSURE



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Figure 14

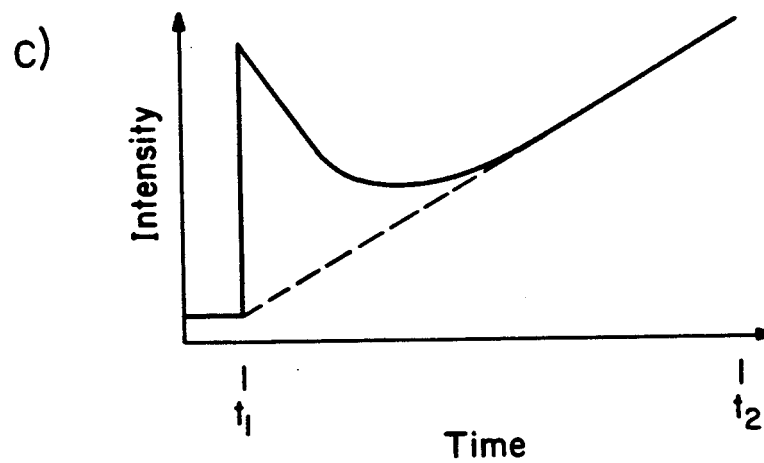
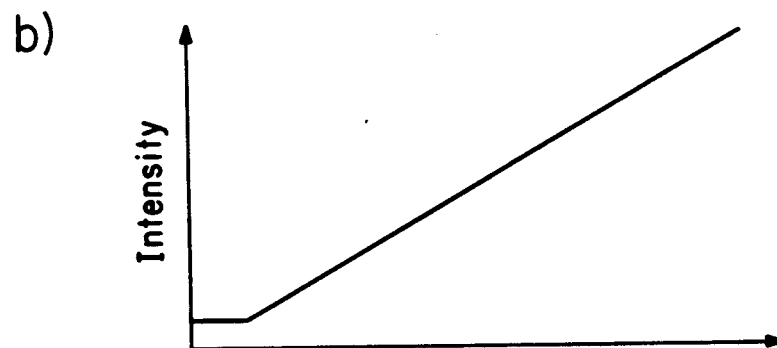
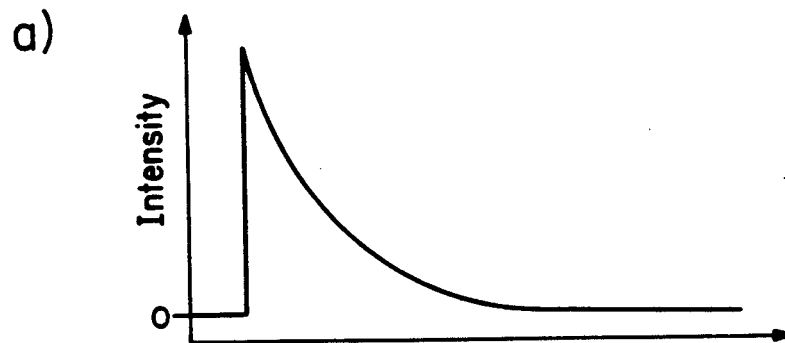


Figure 15

Solidified Zone Thickness vs.
Time After Formation of Microfault

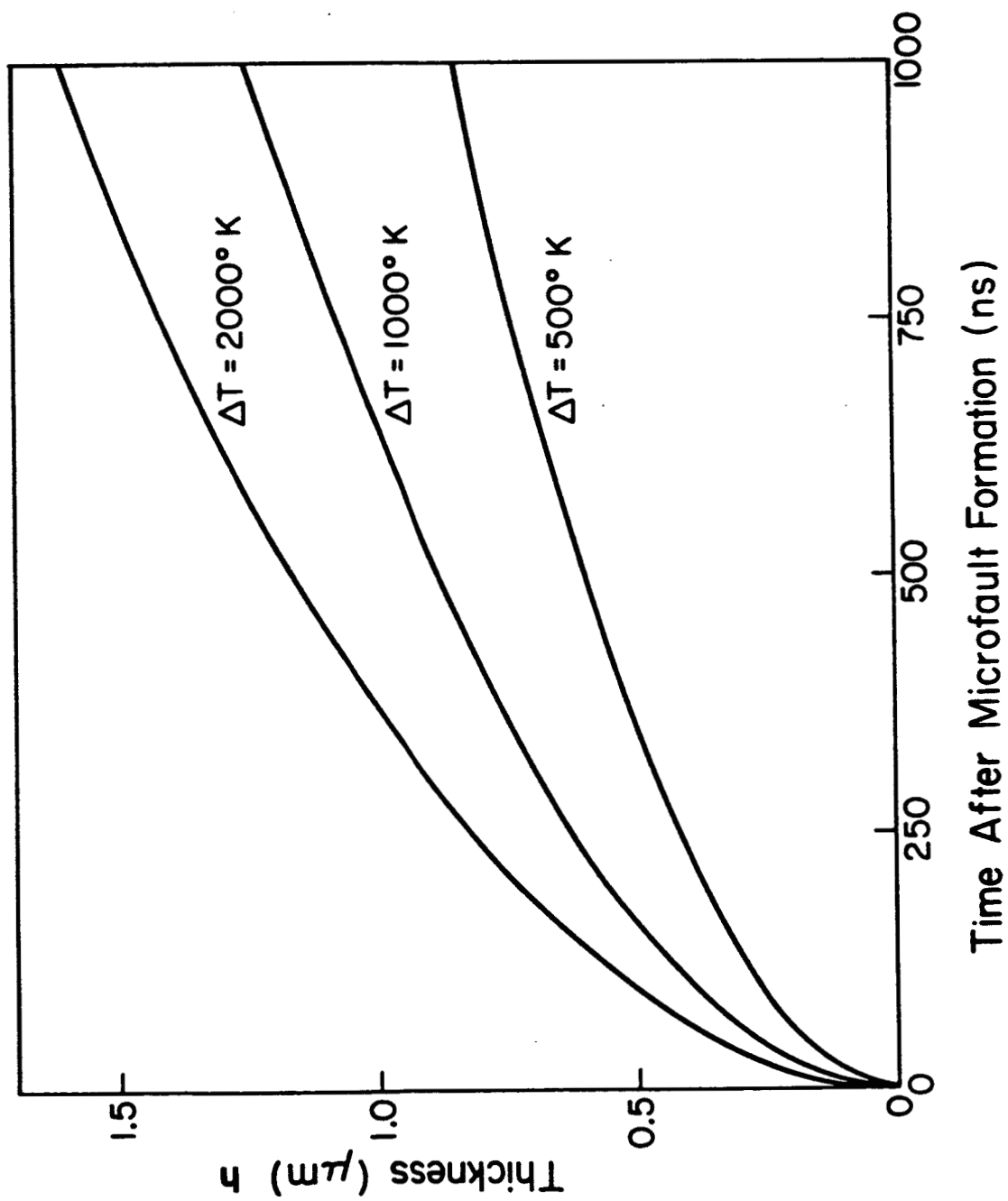


Figure 1b a

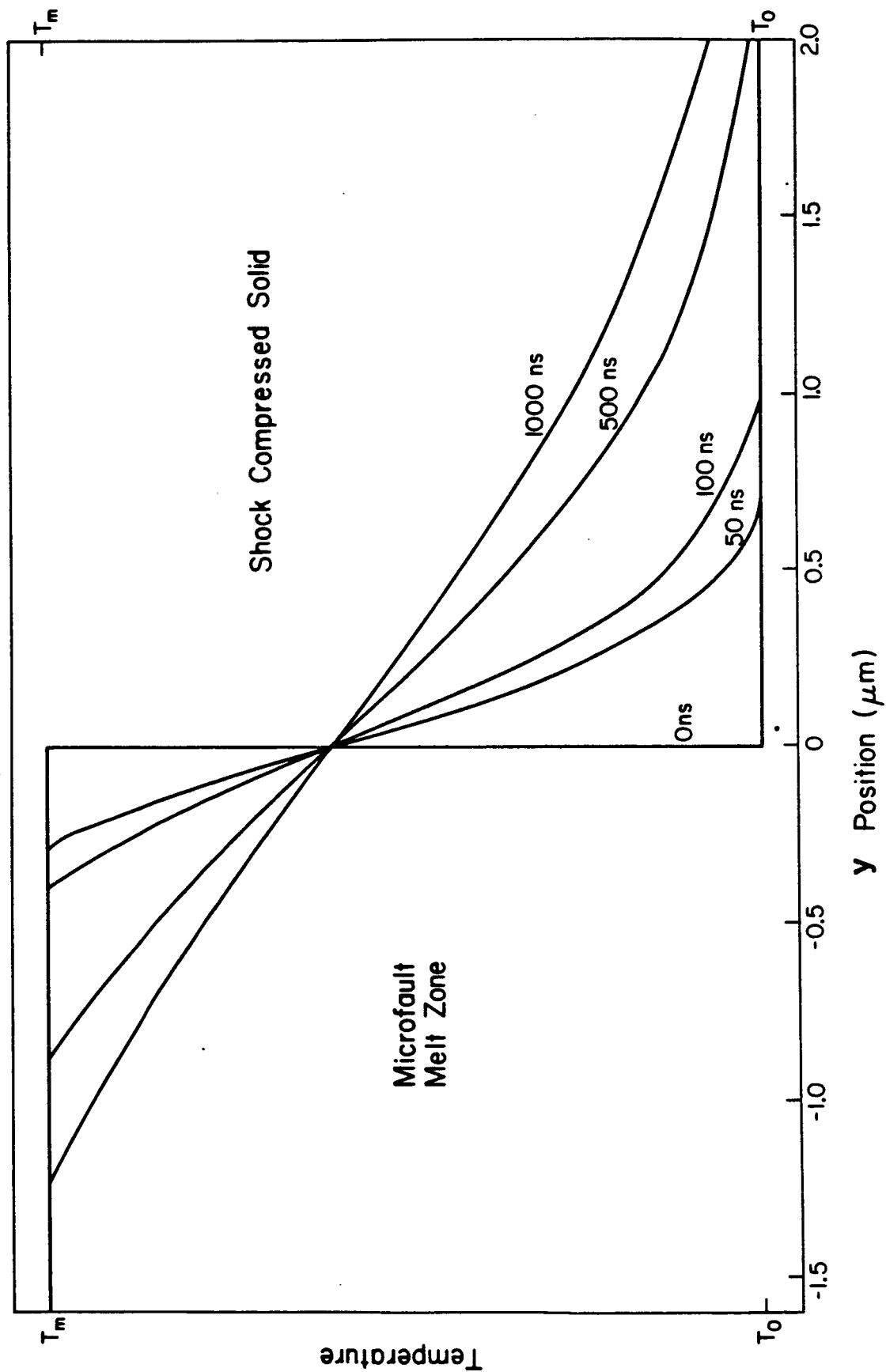
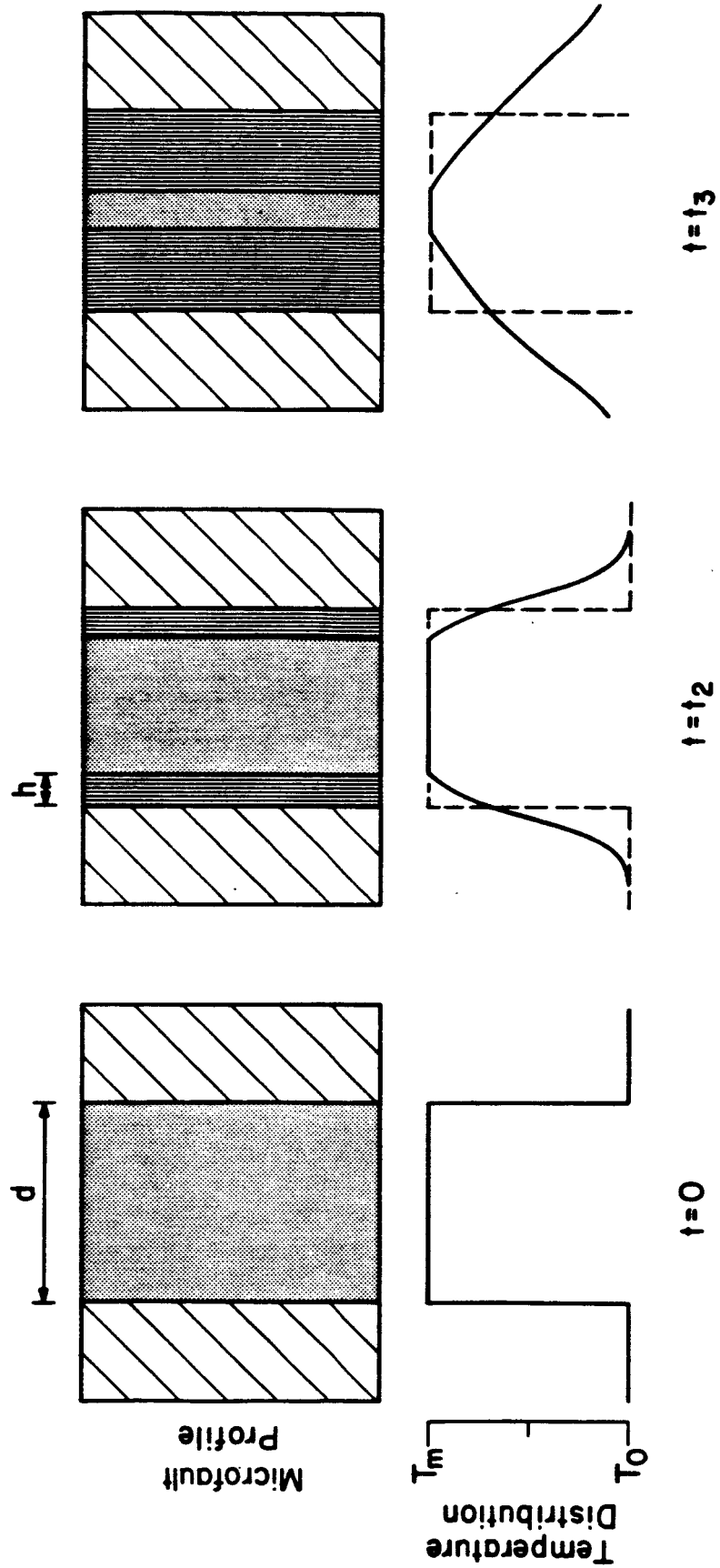
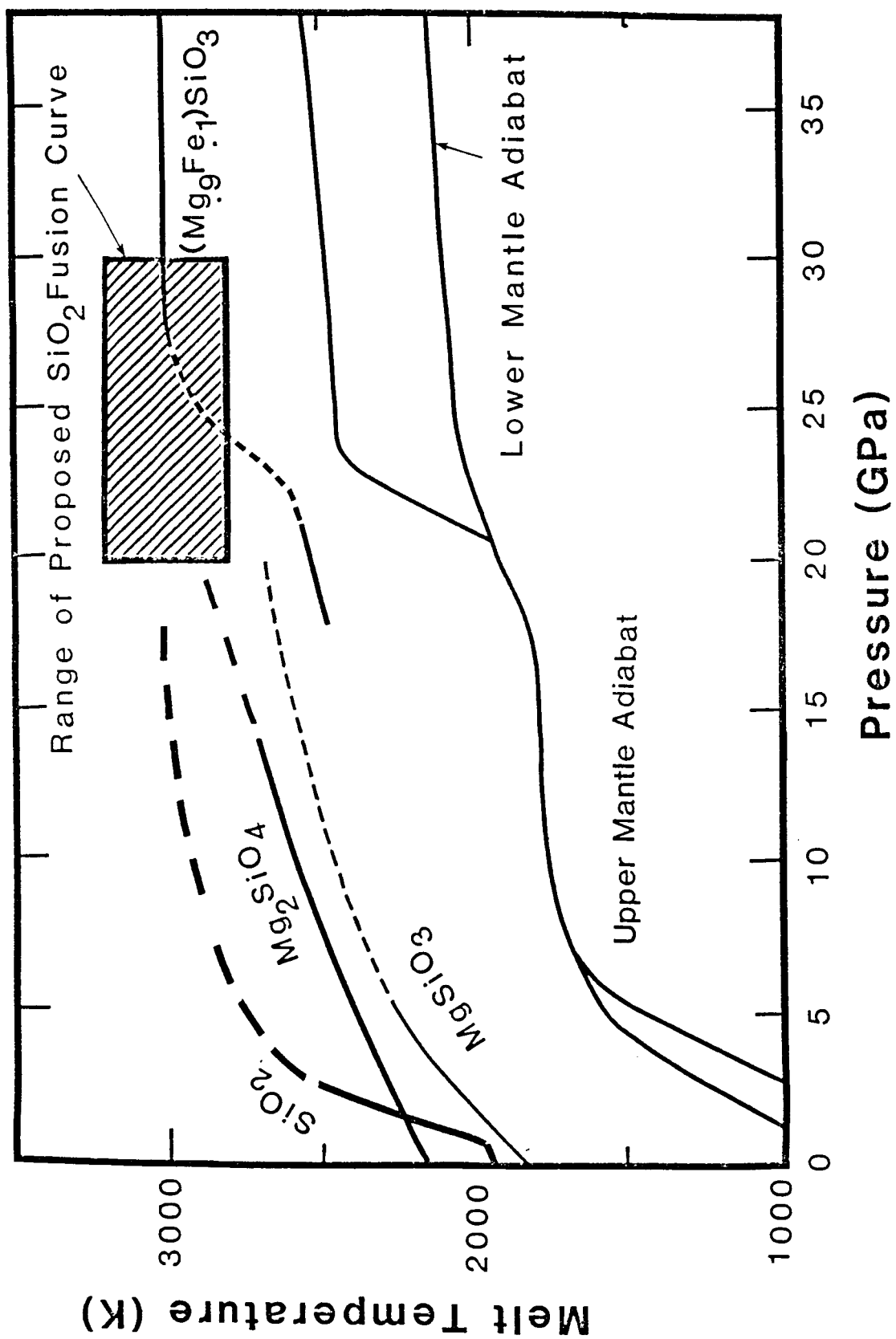
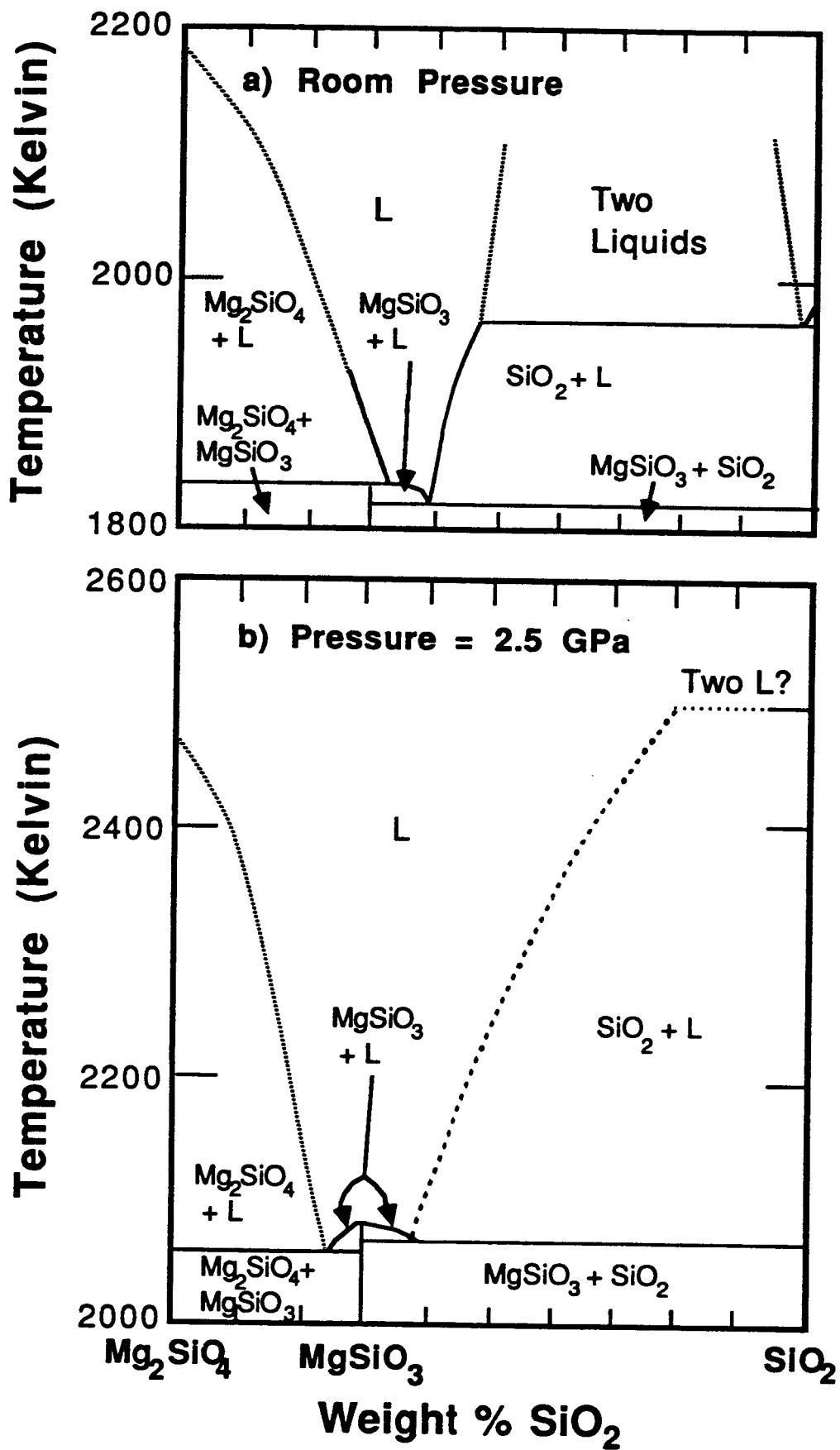
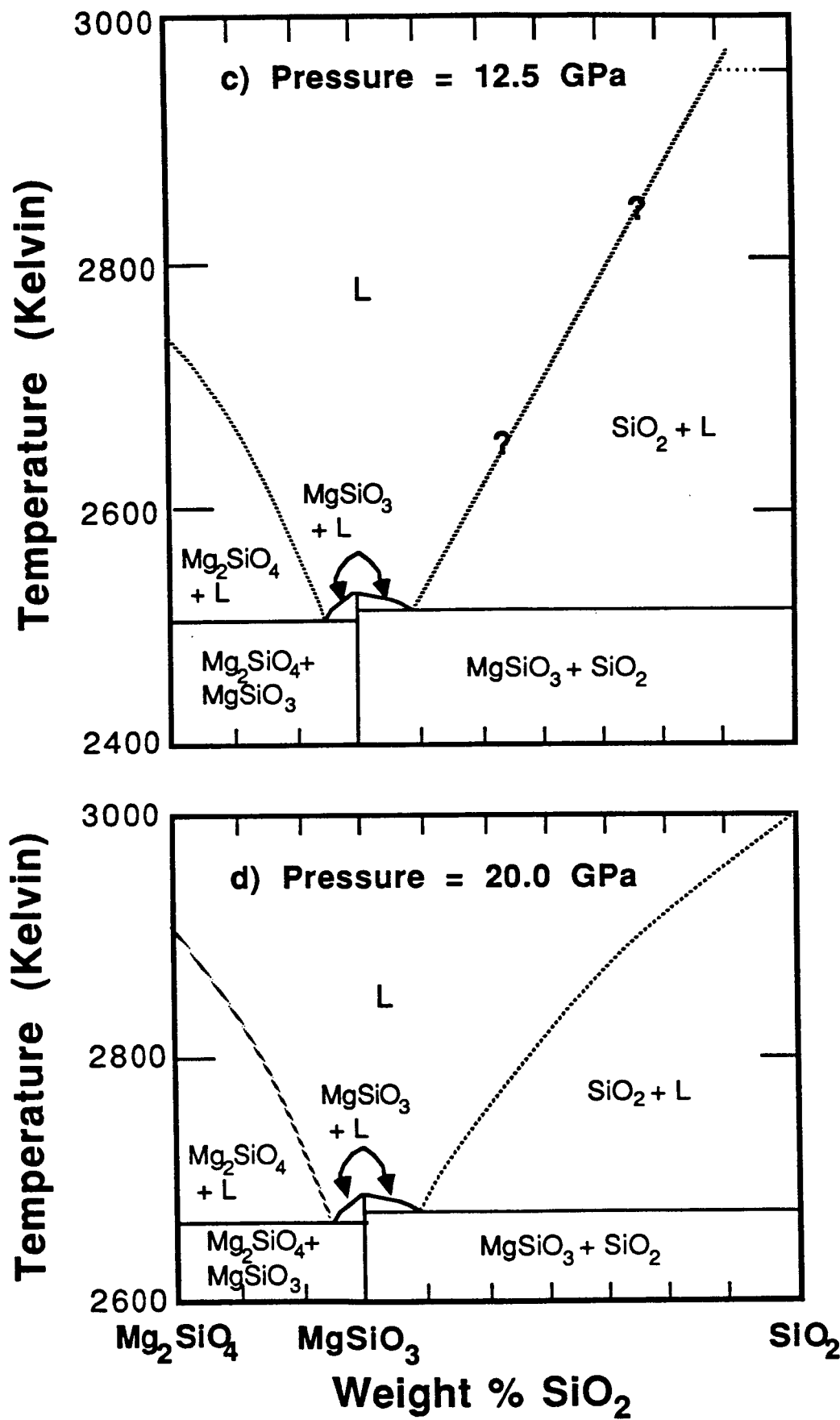


Figure 16b.









TJA87272SC

Figure 19c,d